

Geology and Geophysics Graduate Symposium

Time	Speaker	Title
9:00 am	Annemarie Pickersgill	Shatter cones: A diagnostic feature of hypervelocity impact.
9:12 am	Tanya Harrison	Geologic activity on Enceladus.
9:24 am	Anna Nuhn	Martian, Lunar, and terrestrial cave entrances: A comparative geological analysis.
9:36 am	Eric Pilles	The mantle plume paradigm.
9:48 am	Mahadia Ibrahim	Asteroid mining: Possibilities and challenges.
10:00 am	Cassandra Marion	Volcanism on Io.
10:12 am	Patrick Shepherd	Common causes of colour in natural diamonds: Dislocations, impurities, and vacancies.
10:24 am	Renata Smoke	Formational models of Au-rich volcanogenic massive sulfide deposits.
10:36 am	----- <i>Coffee Break</i> -----	
10:48 am	Yonghua Cao	Convergent margin-related orogenic belts and metallogenesis.
11:00 am	Randy Campbell	Carbonatites: A classification and evolutionary review.
11:12 am	Gloria Eboremen	Bioremediation of petroleum hydrocarbon-contaminated soils: Land farming approach.
11:24 am	David Olutusin	Formation of black shales: Deep versus shallow water interpretation.
11:36 am	Colin Terry	Determining igneous structures and geomagnetic field history from induced and remnant magnetic fields.
11:48 am	Tola Ogunniyi	Resource potential of the Barnett Shale, Fort Worth Basin, Texas.
12:00 noon	Wesley Greig	The precession of the perihelion of Mercury.
12:12 pm	Mansour Al-hashim	Stromatolites: Utility, application, and challenges.
12:24 pm	----- <i>Lunch Break</i> -----	
1:24 pm	Roderick Tom-Ying	Microfossils and the origins of life on Earth.

1:36 pm	Kathryn Lapenskie	Affects of Early to Middle Ordovician climate, paleogeography, and environment on faunal radiation.
1:48 pm	Elizabeth Hooper	Recognition of tsunami deposits within the sedimentary record: Attempts at establishing a tsunami facies.
2:00 pm	Filippo Resente	A comparison of two projects for the prevention of high water in Venice, Italy, as a result of land subsidence and climate-induced sea level rise.
2:12 pm	Tara Despault	Effects of climate-induced temperature and water table changes on carbon dynamics of northern peatlands.
2:24 am	James Goacher	Millennial climate cycles in the Holocene.
2:36 pm	Mengmeng Qu	Strength of the continental lithosphere.
2:48 pm	Weiyin Chen	Shale gas in Canada: Geological controls and current challenges.
3:00 pm	----- <i>Coffee Break</i> -----	
3:12 pm	Alana Crump	Groundwater remediation using zero-valent iron as a reactive medium in permeable reactive barriers.
3:24 pm	Sean Fulcher	Mining salt, brine and clay: A review of lithium and boron evaporite deposits.
3:36 pm	Xiaoming Zhang	Self-organized criticality: what can it tell us about natural hazards?
3:48 pm	Tararat Lerkwieng	Basin controls on the occurrence of reservoir intervals in the Cardium Formation, Alberta.
4:00 pm	Yelena Kropivniskaya	Seismic risk in Canada.
4:12 pm	Behzad Hassani	Uses and challenges in real-time seismological data applications.
4:24 pm	Jonathon Hey	Detachment faulting and it's implications of the mineralization of oceanic core complexes.
4:36 pm	Sean Funk	Models and timing of core formation.
4:48 pm	----- Thank you to everyone for attending! -----	

Please join us for four more presentations on Monday November 26, 9:30 am in BGS 1053

Sarah Sweeney	Dominant gliding versus pure spreading in passive margins: The effect of differential sedimentation on initiating salt tectonics.
Martin Arce	Weathering-induced metal-enrichment processes.
Wajahat Ali	Geophysical techniques for shallow subsurface Ground Penetrating Radar (GPR), Multi-channel Analysis of Shear Waves (MASW), seismic refraction and reflection.
Hadis Samadi Alinia	Database for flooding susceptibility, hazard, and vulnerability assessment.

Shatter Cones: A Diagnostic Feature of Hypervelocity Impact

Annemarie E. Pickersgill

Shatter cones have been an important tool in the identification and study of impact structures on Earth. Most terrestrial impact structures have been heavily eroded, as a result the features which make them readily identifiable on other rocky planetary bodies disappear. This, in addition to the propensity of circular structures of endogenic origin has made the identification of terrestrial impact structures difficult. It is therefore useful to have a unique indicator of shock that is readily identifiable in the field. Shatter cones are the only macroscopic feature that is indicative of shock deformation and therefore diagnostic of hypervelocity impact. They form in large volumes of target rock, and at depth, so they are widespread and often still visible after erosion of the upper part of the structure. An impact origin has been confirmed based on the presence of shatter cones alone, but more often the discovery of shatter cones is followed by a search for microscopic shock metamorphic effects. Shatter cones are found only at impact structures and nuclear test sites, and until recently, only on Earth.

Shatter cones are roughly conical, curved, pervasive fractures characterized by multiple sets of striations that radiate and branch away from the apex. The acute angle of intersection of the striations tends to point toward the apex of the cone. Partial cones are more common than full cones, and the apex of a cone is rarely seen. Smaller “parasitic” cones formed on the surface of larger cones are common, creating a composite texture. Shatter cones range in size from several millimetres to metres. *In situ* shatter cones have been found individually but are far more common in groups, often with roughly parallel axes and with apices pointing in a similar direction. The general direction of orientation is “inward and upward” when beds are restored to pre-impact position. However, cones with highly variable orientations have also been observed in outcrop and hand specimen.

Shatter cones are best developed in fine-grained lithologies, and poorly developed in coarser grained rocks. Crude shatter cones are flatter, and have larger striations that can easily be mistaken for other features such as slickensides, cone-in-cone, wind abrasion features, and anthropogenic blast cones. The most obvious differentiating feature is the penetrative nature of the fractures – if you break a shatter cone it will tend to fracture along other shatter cone surfaces. Microscopic shock metamorphic effects such as planar deformation features and diaplectic glass have been documented in shatter cones, though their presence is not ubiquitous. Similarly localized melting along shatter cone surfaces has been found in some, but not all, samples.

The formation of shatter cones is still poorly understood. Target lithology does not seem to have a large effect on whether or not shatter cones form, only on their quality. Models, experiments, and field studies indicate that shatter cones form immediately as the shock wave passes, at relatively low shock pressures (~2-10 GPa, rarely up to 30 GPa), and prior to excavation of the cavity.

References

- Baratoux D, Melosh HJ. 2003. The formation of shatter cones by shock wave interference during impacting. *Earth and Planetary Science Letters* 216 : 43–54.
- Dietz RS. 1960. Meteorite impact suggested by shatter cones in rock. *Science* 131 : 1781–1784.
- Dietz RS. 1971b. Shatter cones (shock fractures) in astroblemes. *Meteoritics* 6 : 258–259.
- Dressler BO, Sharpton VL. 1997. Breccia formation at a complex impact crater; Slate Islands, Lake Superior, Ontario, Canada. *Tectonophysics* 275 : 285–311.
- Fackelman SP, Morrow JR, Koeberl C, McElvain TH. 2008. Shatter cone and microscopic shock-alteration evidence for a post-Paleoproterozoic terrestrial impact structure near Santa Fe, New Mexico, USA. *Earth and Planetary Science Letters* 270 : 290–299.
- Ferriere L, Lubala FRT, Osinski GR, Kaseti PK, Anonymous. 2011. The Luizi Structure (Democratic Republic of Congo); first confirmed meteorite impact crater in Central Africa. *Abstracts of Papers Submitted to the Lunar and Planetary Science Conference* 42 : 0–1637.
- Ferriere L, Osinski GR, Anonymous. 2010a. Shatter cones and associated shock-induced microdeformations in minerals; new investigations and implications for their formation. *Abstracts of Papers Submitted to the Lunar and Planetary Science Conference* 41 : 0–Abstract 1392.
- Ferriere L, Raiskila S, Osinski GR, Pesonen LJ, Lehtinen M, Anonymous. 2010b. The Keurusselka Structure (Finland); impact origin confirmed by universal-stage characterization of planar deformation features in quartz grains. *Abstracts of Papers Submitted to the Lunar and Planetary Science Conference* 41 : 0–Abstract 1072.
- French BM. 1998. *Traces of catastrophe, a handbook of shock-metamorphic effects in terrestrial meteorite impact structures*. Lunar and Planetary Institute, Houston, TX: Houston, TX, United States (USA)
- French BM, Koeberl C. 2010. The convincing identification of terrestrial meteorite impact structures: What works, what doesn't, and why. *Earth - Science Reviews* 98 : 123.
- French BM, Short NM (ed). 1968. *Shock metamorphism of natural materials; proceedings of the First Conference held at NASA, Goddard Space Flight Center, Greenbelt, Maryland, April 14-16, 1966*
- Gash PJS. 1971. A dynamic mechanism for the formation of shatter cones. *Meteoritics* 6 : 273.
- Huson S, Foit FF, Pope MC, Anonymous. 2006. X-ray diffraction study at Sierra Madera impact structure, West Texas. *Abstracts with Programs - Geological Society of America* 38 : 81.
- Johnson GP, Talbot RJ. n.d. *A theoretical study of the shock wave origin of shatter cones*, 92 pp.
- McHone JF, Shoemaker C, Killgore M, Killgore K, Anonymous. 2012. Two shatter-coned NWA Meteorites. *Abstracts of Papers Submitted to the Lunar and Planetary Science Conference* 43 : 0–Abstract 2359.
- Milton DJ. 1977. Shatter cones; an outstanding problem in shock mechanics. Roddy DJ, Pepin RO, and Merrill RB (eds). Pergamon Press, New York, N.Y.: New York, N.Y., United States (USA)
- Roddy DJ, Davis LK. 1977. *Shatter cones formed in large-scale experimental explosion craters*. Roddy DJ, Pepin RO, and Merrill RB (eds). Pergamon Press, New York, N.Y.: New York, N.Y., United States (USA)
- Sharpton VL, Dressler BO, Herrick RR, Schnieders B, Scott J. 1996. New constraints on the Slate Islands impact structure, Ontario, Canada. *Geology (Boulder)* 24: 851–854.
- Wieland F, Reimold WU, Gibson RL. 2006. New observations on shatter cones in the Vredefort impact structure, South Africa, and evaluation of current hypotheses for shatter cone formation. *Meteoritics & Planetary Science* 41 : 1737–1759.

The Geomorphology of Enceladus

Tanya N. Harrison

Saturn's moon Enceladus, while small enough to fit within the state of Colorado, has garnered the attention of the astronomical and planetary science communities even before any images of the body had been acquired. Ground-based telescopic observations revealed an unusually bright body orbiting within the brightest portion of Saturn's diffuse E ring, leading to the suggestion that Enceladus was a primary source of E ring material. However, the mechanism by which Enceladus could be contributing enough material to account for the observed brightness was unknown. The arrival of the Voyager 1 and 2 spacecraft to the Saturnian system brought about some new understanding of Enceladus with the return of the first photographs of its surface. These photos showed a striking dichotomy, with an older, heavily cratered northern hemisphere and a younger, relatively crater-free southern hemisphere cut by multiple tectonic features. The formation of the tectonic features was attributed to tidal forcing from Saturn and a 2:1 resonance with Dione leading to compressional and extensional stresses. The crater density in the smooth southern region implies that it is nearly the youngest of all the icy satellites in the solar system, third only behind Europa and potentially the polar regions of Triton. How this region was being resurfaced was still unknown from Voyager data. It was not until the arrival of the Cassini mission that the questions of the mechanisms for both resurfacing and E ring contributions would be answered. Cassini observed multiple water ice plumes being ejected from high-temperature (180 K or more in some cases) regions associated with "tiger stripe" fissures in the south polar region. This paper details the progression in our knowledge of the geomorphology Enceladus from the pre-Voyager era to the revolution brought about by Cassini.

References

- Brown, R. H., Clark, R. N., Buratti, B.J., Cruikshank, D.P., Barnes, J.W., Mastrapa, R.M.E, Bauer, J., Newman, S., Momary, T., Baines, K.H., Bellucci, G., Capaccioni, F., Cerroni, P., Combes, M., Coradini, A., Drossart, P., Formisano, V., Jaumann, R., Langevin, Y., Matson, D.L., McCord, T.B., Nelson, R.M., Nicholson, P.D., Sicardy, B., and Sotin, C. 2006. Composition and physical properties of Enceladus' surface. *Science*, **311**, 1425–1428. doi:10.1126/science.1121031.
- Hansen, C.J., Esposito, L., Stewart, A.I.F., Colwell, J., Hendrix, A., Pryor, W., Shemansky, D., and West, R. 2006. Enceladus' water vapor plume. *Science*, **311**, 1422–1425. doi:10.1126/science.1121254.
- Kargel, J.S., and Pozio, S. 1996. The volcanic and tectonic history of Enceladus. *Icarus*, **119**, 385–404.
- Porco, C.C., Helfenstein, P., Thomas, P.C., Ingersoll, A.P., Wisdom, J., West, R., Neukum, G., Denk, T., Wagner, R., Roatsch, T., Kieffer, S., Turtle, E., McEwen, A., Johnson, T.V., Rathburn, J., Veverka, J. Wilson, D., Perry, J., Spitale, J., Brahic, A., Burns, J.A., DelGenio, A.D., Dones, L., Murray, C.D., and Squyres, S. 2006. Cassini observes the active south pole of Enceladus. *Science*, **311**, 1393–1401. doi:10.1126/science.1123013.
- Smith, B.A., Soderblom, L., Batson, R., Bridges, P., Inge, J., Masursky, H., Shoemaker, E., Beebe, R., Boyce, J., Briggs, G., Bunker, A., Collins, S.A., Hansen, C.J., Johnson, T.V., Mitchell, J.L., Terrile, R.J., Cook, A.F., Cuzzi, J., Pollack, J.B., Danielson, G.E., Ingersoll, A.P., Davies, M.E., Hunt, G.E., Morrison, D., Owen, T., Sagan, C., Veverka, J., Strom, R., and Suomi, V.E. 1982. A new look at the Saturn system: The Voyager 2 images. *Science*, **215**, 504–537.
- Spencer, J.R., Pearl, J.C., Segura, M., Flasar, F.M., Mamoutkine, A., Romani, P., Buratti, B.J., Hendrix, A.R., Spilker, L.J., and Lopes, R.M.C. 2006. Cassini encounters Enceladus: Background and the discovery of a south polar hot spot. *Science*, **311**, 1401–1405. doi:10.1126/science.1121661.
- Squyres, S.W., Reynolds, R.T., Cassen, P.M., and Peale, S.J. 1983. The evolution of Enceladus. *Icarus*, **53**, 319–331.

Martian, Lunar, and Terrestrial Cave Entrances: A Comparative Geological Analysis

Anna Nuhn

Large-scale igneous provinces are found on both the Martian and Lunar landscapes which make for excellent environments for near surface basaltic lava tubes and associated atypical pit crater formations. Models of lava tube formation typically involve the outer surface of a lava channel cooling more rapidly consequently forming a hardened crust; the remaining lava flows out of the tube, leaving a void space. Models for atypical pit crater formation include collapsed lava tubes, dilational faulting, dyke swarms and collapsed magma chambers, all acting as subsurface voids for surface collapse. Near surface basaltic lava tubes and associated atypical pit craters have recently been known to possess cave-like entrances on the Moon and Mars, called “skylights”.

Caves on Mars and the Moon have been hypothesized since the late 1960s, but not until recently these planetary structures were discovered. This paper will review newly observed lunar skylights in Marius Hills, Mare Tranquillitatis, and Mare Ingenii along with Martian skylights in the flanks of Pavonis, Ascraeus, and Arsia Mons, and flows from Hadriaca Patera. Observations of these entrances have been done using a suite of improved orbital high-resolution imagery and thermal infrared detection methods of the subsurface geology. For Mars these instruments include, the Mars Odyssey’s Thermal Emission Imaging System (THEMIS), the Mars Reconnaissance Orbiter’s Context Camera (CTX) and High Resolution Imaging System (HiRISE). On the Moon, KAGUYA SELEnological and Engineering Explorers’(SELENE) Terrain Camera (TC) and the Multi-band Imager (MI), as well as the Lunar Reconnaissance Orbiter Camera (LROC) Narrow Angle Camera (NAC) are being utilized.

These extraterrestrial cave entrances pose as ideal spaces for future robotic exploration, mineral and astrobiology investigations, as well as potential human bases in planetary missions. Understanding cave entrance formation on Earth is critical for understanding the development and exploration possibilities on other planetary bodies. Terrestrial analogue sites including the Atacama Desert in northern Chile, Teide National Park in Canary Islands Tenerife, Spain, and Kazumura Cave in Kilauea, Hawaii, will continue to serve as locations for the development of scientific exploration strategies and new technologies for future Martian and Lunar missions.

References

- Cushing, G.E., Titus, T.N., Wynne, J.J., and Christensen, P.R. 2007. THEMIS observes possible cave skylights on Mars [online] *Geophysical Research Letters*, **34**: L17201. doi:10.1029/2007GL030709.
- Cushing, G. E. 2012. Candidate cave entrances on Mars [online]. *Journal of Cave and Karst Studies*, **74** (1): 33–47. doi: 10.4311/ 2010EX0167R.
- Greeley, R. 1971a, Observations of actively forming lava tubes and associated structures. Hawaii: *Modern Geology*, **2**: 207–223.
- Greeley, R. 1971b. Lava tubes and channels in the Lunar Marius Hills [online]. *Earth, Moon, and Planets*, **3**(3): 289–314. doi:10.1007/BF00561842.
- Haruyama, J., Hioki, K., Shirao, M., Morota, T., Hiesinger, H., van der Bogert, C.H., Miyamoto, H., Iwasaki, A., Yokota, Y., Ohtake, M., Matsunaga, T., Hara, S., Nakanotani, S., Pieters, C.M. 2009b. Possible lunar lava tube skylight observed by SELENE cameras. *Geophysical Research Letters*, **36**: L21206. doi:10.1029/2009GL040635.
- Keszthelyi, L., and Self, S. 1998. Some physical requirements for the emplacement of long basaltic lava flows [online]. *Journal of Geophysical Research*, **103**(B11): 27447-27464. doi:10.1029/98JB00606.
- Leveille, R.J. and Datta, S. 2010. Lava tubes and basaltic caves as astrobiological targets on Earth and Mars: A review. *Planetary and Space Science*, **58**: 592–598. doi:10.1016/j.pss.2009.06.004.
- Miyamoto, H., Haruyama, J., Kobayashi, T., Suzuki, K., Okada, T., Nishibori, T., Showman, A., Lorenz, R., Mogi, K., Crown, D.A., Rodriguez, J.A.P., Rokugawa, S., Tokunaga, T., and Masumoto, K. 2005, Mapping the structure and depth of lava tubes using ground penetrating radar [online]. *Geophysical Research Letters*, **32**: L21316. doi:10.1029/2005GL024159.
- Morse, A.D., Laines A., and Howard, K.T. 2011. Exploration of lava tubes in the Teide National Park, a martian analog. *In* First International Planetary Caves Workshop: Implications for Astrobiology, Climate, Detection, and Exploration, Lunar and Planetary Institute, Houston, LPI Contribution No. 1640, pp. 22-23.
- Oberbeck, V.R., Quaide, W.L., and Greeley, R. 1969. On the origin of lunar sinuous rilles. *Mod. Geol.* **1**: 75–80.
- Valerio, A., Tallarico, A., Dragoni, M. 2008. Mechanisms of formation of lava tubes [online]. *Geophysical Research Letters*, **113**: B08209. doi:10.1029/2007JB005435.
- Wynne, J.J., Titus, T.N., and Chong Diaz, G. 2008. On developing thermal cave detection techniques for Earth, the Moon and Mars [online]. *Earth and Planetary Science Letters*. **272**: 240–250. doi:10.1016/j.epsl.2008.04.037.
- Wyrrick, D., D. A. Ferrill, A. P. Morris, S. L. Colton, and D. W. Sims. 2004. Distribution, morphology, and origins of Martian pit crater chains [online]. *Journal of Geophysical Research Letters*, **109**: E06005. doi:10.1029/2004JE002240.

The Mantle Plume Paradigm

Eric Pilles

Modern mantle plume theory is incomplete. Numerous publications dispute portions of the theory, while others have raised the question if mantle plumes exist at all. Alternative theories have been produced which directly relate to plate-tectonic mechanisms and completely disregard mantle plume theory. When considering the evidence, from both sides of the argument, it is clear that while the current mantle plume model is incomplete, it is superior to alternative theories presented by the “Antiplume Lobby”.

The current mantle plume model defines mantle plumes as a spatially fixed upwelling of hot light material that ascends from the core-mantle boundary. The theory states that the plumes themselves are stationary while the plates above them move freely.

There are several objections to the mantle plume theory. First, geochemical indicators such as $^3\text{He}/^4\text{He}$ are often used to indicate origin from the mantle-core interface. However, large-scale mixing of the lithosphere and mantle - via subduction and convection - result in chemical heterogeneity of the mantle. Mantle plumes are not always associated with an uplift of the Earth's surface, for example at the Siberian flood basalt province, the submarine Ontong Java Plateau, and the Decan traps. The statement that plumes are stationary has been proven false. Displacement has been observed in both the head and tails of plumes. Additionally, the anomalously high temperatures necessary for melting under dry conditions would be accompanied by increased heat flow above the plume, however this is not seen.

Alternative theories suggest that the ‘plume’ feature is attributed to plate tectonic mechanisms and appear as a result of shallow tectonic stress, subsequent decompression, and melting of the mantle enriched in basaltic material. However, these theories fail to address two major problems – problems which the mantle plume theory can explain. Time-progressive volcanic chains, such as those at Hawaii, a classic example of mantle plumes, cannot be explained by alternative theories such as crack propagation – while mantle plume theory explains not only the general behaviour of volcanic chains, but also many details related to plate motion in these regions. Additionally, alternative theories involve passive plate tectonic mechanisms, which take place in the lithosphere and upper asthenosphere, while seismic tomographic data has shown that plumes can extend into the mesosphere. Therefore, while many answers remain unanswered, the mantle plume theory still remains the most accurate model to date.

References

- Balyshev S.O. and Ivanov, A.V. 2001. *Low-Density Anomalies in the Mantle: Ascending Plumes and/or Heated Fossil Lithospheric Plates?*, Dokl. Akad. Nauk, **380**: Issue 4, 523–527.
- Camp, V.E., 1995. *Mid-Miocene propagation of the Yellowstone mantle plume head beneath the Columbia River Basalt source region*, *Geology*, **23**: 435–438, doi:10.1130/00917613(1995)023<0435:MMPOTY>2.3.CO;2.
- Christiansen, R., Foulger, G., and Evans, J., 2002. *Upper-mantle origin of the Yellowstone hotspot*, *Geological Society Of America Bulletin*, **114**: 1245–1256, doi:10.1130/0016-7606.
- Ernst, R. and Buchan, K., 2003. *Recognizing Mantle Plumes in the Geological Record*, *Ann. Rev. Earth Planet. Sci.* **31**: 459–523.
- Fouch, M. 2012. *The Yellowstone Hotspot: Plume or Not?*, *Geology*, **40**: 479–480. Doi: 10.1130/focus052012.1.
- Gorbatov, A., Fukao, Y., Widiyantoro, S. and Gordeev, E., 2001. *Seismic Evidence for a Mantle Plume Oceanwards of the Kamchatka–Aleutian Trench Junction*, *Geophys. J. Int.* **146**: 282–288.
- Graham, D.W., Reid, M.R., Jordan, B.T., Grunder, A.L., Leeman, W.P., and Lupton, J.E., 2009. *Mantle source provinces beneath the northwestern USA delimited by helium isotopes in young basalts*, *Journal of Volcanology and Geothermal Research*, **188**: 128–140, doi:10.1016/j.jvolgeores.2008.12.004.
- Humphreys, E., Dueker, K., Schutt, D., and Smith, R., 2000. *Beneath Yellowstone: Evaluating plume and nonplume models using teleseismic images of the upper mantle*, *GSA Today*, **10**: 1–7.
- Morgan, W.J., 1971. *Convection plumes in the lower mantle*, *Nature*, **230**: 42–43, doi:10.1038/230042a0.
- Puchkov, V. 2009. *The Controversy over Plumes: Whi Is Actually Right?*, *Geotectonics*, **43**: No. 1, 1-17.
- Sheth, H.C., 2007. *Plume-Related Regional Pre-Volcanic Uplift in the Deccan Traps: Absence of Evidence, Evidence of Absence*, in *Plates, Plumes, and Planetary Processes*, Ed. by G. R. Foulger and D. M. Jurdy (Geol. Soc. Am. Spec. Paper **430**, 2007), 785–814 (2007).
- Wolfe, C.J., Solomon, S.C., Laske, G., Collins, J.A., Detrick, R.S., Orcutt, J.A., Bercovici, D., and Hauri, E.H., 2009. *Mantle shear-wave velocity structure beneath the Hawaiian hot spot*, *Science*, **326**: 1388– 1390, doi:10.1126/science.1180165.
- D. Zhao, 2004. *Global Tomographic Images of Mantle Plumes and Subducting Slabs: Insight into Deep Earth Dynamics*, *Phys. Earth Planet. Inter.* **146**, 3–34.

Asteroid Mining: Possibilities and Challenges

Mahdia Ibrahim

Asteroids represent the remaining building blocks from the early Solar System formation ~ 4.6 billion years ago. Most asteroids have retained a relatively pristine record of nebular and early planet-forming processes that provides clues about Solar System dynamics and orbital evolution. Improved remote observations, the science of meteoritics, and multiple flyby and rendezvous missions have provided much of our knowledge on asteroid regoliths, their mineralogical composition and physical properties. Detailed analysis of a number of meteorite samples showed relatively high concentrations of valuable metals by Earth standards; including platinum group elements, Rare Earth Elements (REE) and gold. Asteroid mining may lead to economic gains and scientific outcomes that will contribute greatly to the state of human civilization on Earth and beyond.

The idea of the possible exploitation of asteroidal resources of minerals and REE is not new (early 1900's), but substantial data about the composition and surface properties of asteroids were unavailable at the time. Today, science data and technology necessary for asteroid mining are becoming available. Assuming that enough knowledge has been acquired on asteroids' properties, mining remains a challenge for many reasons that could be summarized in three areas: 1) economic feasibility and market demand; 2) (NEA) target selection and orbital dynamics; and, 3) mission design and operations.

While successful mining of asteroids have the advantage of obtaining resources without losing energy in huge gravity wells such as on the moon, the cost of mining might still outweigh the desired economic outcome. This has been addressed by suggestions of the utilization of products in low-Earth-orbit (LEO); on lunar bases or future space stations, which could provide an excellent market as contractors and operators in LEO will utilize the material and cut down further costs of transportation. From a science and engineering perspective, a suitable target would ideally be from the near-Earth Asteroid (NEA) population with low inclination to reduce the costs of launch and recovery. Moreover, asteroids are dynamic and extreme environments in terms of temperature, radiation, and physical characterization (e.g. rubble pile asteroids), which requires creative and unconventional technology different than what is known in terrestrial mining. Mission design will require anchoring securely to a "moving object, extracting material depending on lithology (fragmentation on silicate lithologies versus vaporization on hydrated lithology), followed by the challenge of material storage and transportation. This leads to suggestions of robotic capture and retrieval of asteroids to near-Earth orbit, where in-situ utilization can be conducted, thus efficiently cutting down the cost of mining and gaining control of near-Earth objects (NEOs) for multiple purposes including development of asteroid deflection technology.

References

- Brophy, J., and 32 authors. 2012. *Asteroid Retrieval Feasibility Study*. Keck Institute for Space Studies, California Institute of Technology, Jet Propulsion Laboratory. Internal report series.
- Blair, B. R., 2000. *The Role of Near-Earth Asteroids in Long-Term Platinum Supply*. Space Resources roundtable 2, pp.5
- Elvis, M. 2012. *Let's mine asteroids – for science and profit*. Nature., **485**: 549
- Forgan, D. H., and Elvis, M. 2011. *Extraterrestrial asteroid mining as forensic evidence for extraterrestrial intelligence*. *International Journal of Astrobiology*. **10 (4)**: 307-313
- Gerlach, C. L. 2005. *Profitably exploiting near-Earth object resources*, 2005 International Space Development Conference. National Space Society, Washington, DC, May 19-22
- Kargel, J. S. 1994. *Metalliferous asteroid as potential sources of potential metals*, J. Geo. Res., **99**: No. E10, 21,129-21,141
- Metzger, P. T. 2011. *Nature's Way of Making Audacious Space Projects Viable*. DARPA 100 Year Starship Symposium, Orlando, Florida
- Nichols, C. R. 1993. *Volatile products from carbonaceous asteroids*. J. Lewis, M.S. Matthews, M.L. Guerrier (Eds.), Resources of Near-Earth Space, University of Arizona Press, Tucson, Arizona
- O'Leary, B. 1977. *Mining the Apollo and Amor Asteroids*. Science, **197**: No. 4301. pp. 363-366
- Ross, S. D. 2001. *Near-Earth Asteroid Mining*. Space Industry Report. Internal report: Control and dynamical Systems, Caltech, Pasadena, CA.
- Sonter, M. J. 1997. *The technical and economic feasibility of mining the near-Earth asteroids*. Acta Astronautica. **41**: No. 4 – 10, pp. 637 – 647
- Sonter, M. J. 2001. *Near-Earth objects as resources for space industrialization*. Solar System Development Journal **1(1)**: 1-31

Volcanism on Io

Cassandra L. Marion

This manuscript examines the current state of knowledge of volcanism on Io. Io is Jupiter's closest Galilean satellite and the most volcanically active planetary body in the solar system. Unlike all other planetary bodies in the solar system, it lacks evidence of impact cratering, indicating a young surface. Remote sensing techniques applied through the use of ground-based observations and fly-by missions, such as Voyager and Galileo, have recorded spectacular images and spectral data that have led to exciting discoveries. Io's global heat output is estimated to be 25 times greater than Earth, at 10^{14} W. It is unique in the solar system in that its primary internal heat source is tidal heating. Due to its Laplace resonance with neighbouring moons, Ganymede and Europa, energy is dissipated internally, melting a large amount of the interior. The extent of melting and mechanisms of heat transfer within Io are uncertain, however based on its bulk density studies have shown that Io is a differentiated body and likely has an iron or iron sulfide core.

Eruption styles on Io range from flow-dominated to explosion-dominated to intra-patera volcanism. These occur in the form of lava flow fields, lava fountains and lava lakes to explosive, volatile-driven, umbrella-shaped plumes of gas and dust ejected several hundred metres high. Io's tenuous atmosphere is formed primarily by plumes. They occur either in numerous smaller plumes, which are produced near the margins of active lava flows by interactions with near-surface to surface SO₂ ice, or as giant plumes that can reach >200 km high. The dominant volatiles on Io, driving explosive volcanism, are sulfur and sulfur dioxide. There is little evidence of effusive sulfur volcanism, but much of Io is blanketed in SO₂ snow from plume fallout. Eruption temperatures indicate Io's dark lava flows are mafic to ultramafic in composition. However eruption temperatures may not be reflective of liquidus temperatures of the magma, due to either super-heating during magma ascent to the surface, rapid-cooling once extruded, or both. Future missions will further investigate the unknown features and processes on Io.

Selected References

- Cataldo E. 2002. A model for large-scale volcanic plumes on Io: Implications for eruption rates and interactions between magmas and near-surface volatiles. *Journal of Geophysical Research* **107**: 1–12.
- Davies A. 2007. *Volcanism on Io: a comparison with Earth*. Cambridge University Press: New York
- Geissler P. and McMillan M. 2008. Galileo observations of volcanic plumes on Io. *Icarus* **197**: 505–518.
- Keszthelyi L. et al. 2001. Imaging of volcanic activity on Jupiter's moon Io by Galileo during the Galileo Europa Mission and the Galileo Millennium Mission. *Journal of Geophysical Research* **106**: 33025–33052.
- Keszthelyi L., Jaeger W., Milazzo M, Radebaugh J, Davies AG, Mitchell KL. 2007. New estimates for Io eruption temperatures: Implications for the interior. *Icarus* **192**: 491–502.
- Lainey V., Arlot J-E., Karatekin O., Van Hoolst T. 2009. Strong tidal dissipation in Io and Jupiter from astrometric observations. *Nature* **459**: 957–9.
- Lopes-Gautier R, McEwen A. 1999. Active volcanism on Io: Global distribution and variations in activity. *Icarus* **140**: 243–264.
- Matson D., Ransford G., Johnson T. 1981. Heat flow from Io. *Journal of Geophysical Research* **86**: 1664–1672.
- McEwen A.S. et al. 1998. High-Temperature Silicate Volcanism on Jupiter's Moon Io. *Science* **281**: 87–90.
- Morabito L., Synnott S., Kupferman P. and Collins S. 1979. Discovery of Currently Active Extraterrestrial Volcanism. *Science* **204**: 972.
- Peale S., Cassen P., Reynolds R. 1979. Melting of Io by tidal dissipation. *Science* **203**: 892–894.
- Rathbun J. A. and Spencer J.R. 2006. Loki, Io: New ground-based observations and a model describing the change from periodic overturn. *Geophysical Research Letters* **33**: 1-5.
- Veeder G.J., Davies A.G., Matson D.L., Johnson T.V., Williams D. A., Radebaugh J. 2012. Io: Volcanic thermal sources and global heat flow. *Icarus* **219**: 701–722.
- Veeder G.J, Matson D.L., Johnson T.V., Blaney D.L., Goguen J.D. 1994. Io's heat flow from infrared radiometry: 1983–1993. *Journal of Geophysical Research* **99**: 17095.
- Williams D. A., Keszthelyi L.P., Crown D. A., Yff J. A., Jaeger W.L., Schenk PM, Geissler PE, Becker TL. 2011. Volcanism on Io: New insights from global geologic mapping. *Icarus* **214**: 91–112.
- Williams D.A. and Howell R.R. 2007. Active Volcanism: Effusive eruptions. In *Io After Galileo: A new view of Jupiter's volcanic moon*. Springer-Praxis Publishing Ltd.: Germany; 133–161.
- Wilson L. 2009. Volcanism in the Solar System. *Nature Geoscience* **2**: 389–397.
- Wilson L. and Head J.W. 1981. Ascent and eruption of basaltic magma on the Earth and Moon. *Journal of Geophysical Research* **86**: 2971–3001.
- Yoder C. and Peale S. 1981. The tides of Io. *Icarus* **47**: 1–35.

Common causes of colour in natural diamonds: dislocations, impurities, and vacancies

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Diamonds are crystalline carbon, which are extremely valuable as gemstones due to their hardness, fire and brilliance. Diamond gemstones alone generate an annual $\sim \$2 \times 10^{10}$ USD, excluding diamonds used for industrial purposes. The general public typically thinks of diamonds as colourless, while in fact the most valuable diamonds are coloured. Diamonds with desirable colours are referred to as having a “fancy colour”, and are typically deep shades of blue, pink, yellow, purple, green, and red. Due to the value of coloured diamonds, much research has been put into understanding the source of the colour, especially in light of artificially-produced diamonds with high-pressure – high-temperature (HTPT) treatments altering colour. Blue diamonds are both associated with boron impurities in the crystal structure, whereas yellow is caused by nitrogen (N₂ or N₃) or hydrogen. Purple, pink, and red are all associated with dislocations, or vacancies, associated with plastic deformation of the crystal. Green colour can be attributed broadly to two main causes: irradiation by gamma-rays (GR1), or nickel impurities. Brown diamonds have recently become an area of interest, because it is now possible to remove the colouring with HTPT treatments, significantly increasing the value of the gemstones. The general consensus is that the brown colouring is caused by plastic deformation, although this remains controversial. It is also debated whether the brown colour was formed before deposition in the upper mantle or during ascent within the kimberlite pipe. The lack of a consolidated review of the origin of colour in diamonds has left hypotheses difficult to compare to one another.

References

- De Weerd, F. and Van Royen, J. 2001. Defects in coloured natural diamonds. *Diamond and Related Materials*, **10**: 474-479.
- Fisher, D., Sibley, S.J., and Kelly, C.J. 2009. Brown colour in natural diamond and interaction between the brown related and other colour-inducing defects. *Journal of Physics Condensed matter : an Institute of Physics journal*, **21**: 364213.
- Gaillou, E., Post, J., Rost, D., and Butler, J. 2012. Boron in natural type IIb blue diamonds: Chemical and spectroscopic measurements. *American Mineralogist*, **97**: 1-18.
- Gaillou, E., Post, J.E., Bassim, N.D., Zaitsev, A.M., Rose, T., Fries, M.D., Stroud, R.M., Steele, A., and Butler, J.E. 2010. Spectroscopic and microscopic characterizations of color lamellae in natural pink diamonds. *Diamond & Related Materials*, **19**: 1207-1220.
- Godfrey, I.S. and Bangert, U. 2010. Atomic structure-colour relationship in natural diamonds. *Journal of Physics: Conference Series*, **241**: 012053.
- Harlow, G.E. 1997. *The nature of diamonds*. Cambridge University Press.
- Jones, R. 2009. Dislocations, vacancies and the brown colour of CVD and natural diamond. *Diamond & Related Materials*, **18**: 820-826.
- King, J. and Wang, W. 2004. Unusual cause of blue color in a diamond. *GEMS & GEMOLOGY*, **40**: 245-245.
- Kitawaki, H. 2007. Gem diamonds: Causes of colors. *New Diamond And Frontier Carbon Technology*, **17**: 119.
- Massi, L., Fritsch, E., Collins, A.T., Hainschwang, T., and Notari, F. 2005. The "amber centres" and their relation to the brown colour in diamond. *Diamond & Related Materials*, **14**: 1623-1629.
- Smith, E., Helmstaedt, H., and Flemming, R. 2010. Survival of the brown color in diamond during storage in the subcontinental lithospheric mantle. *Canadian Mineralogist*, **48**: 571-582.
- Titkov, S.V., Ivanov, A.I., Marfunin, A.S., Bershov, L.V., Kulakov, V.M., and Chukichev, M.V. 1995. Irradiation as the cause of the bulk green color in natural diamonds. *Transactions Doklady - Russian Academy of Sciences: Earth Science Sections*, **337**: 133-138.
- Titkov, S., Krivovichev, S., and Organova, N. 2012. Plastic deformation of natural diamonds by twinning: Evidence from X-ray diffraction studies. *Mineralogical Magazine*, **76**: 143-149.
- van der Bogert, C.H., Smith, C.P., Hainschwang, T., and McClure, S.F. 2009. Gray-to-blue-to-violet hydrogen-rich diamonds from the argyle mine, australia. *Gems & Gemology*, **45**: 20-37.
- Wang, W., Hall, M., and Breeding, C.M. 2007. Natural type Ia diamond with green-yellow color due to ni-related defects. *Gems & Gemology*, **43**: 240-243.
- Willems, B., Martineau, P.M., Fisher, D., Van Royen, J., and Van Tendeloo, G. 2006. Dislocation distributions in brown diamond. *Physica Status Solidi (a)*, **203**: 3076-3080.

Formational models of Au-rich volcanogenic massive sulfide deposits

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Volcanogenic massive sulfide (VMS) deposits include: volcanic-associated, volcanic-hosted, and volcano-sedimentary hosted massive sulfide deposits; and are major sources of Zn, Cu, Pb, Ag, and Au; and significant sources of Co, Sn, Se, Mn, Cd, In, Bi, Te, Ga, and Ge. Gold-rich VMS deposits form a unique subset of VMS deposits, and, like typical VMS deposits, are found: in sub-marine volcanic terranes with compositions ranging from mafic bi-modal to felsic bi-modal to bi-modal siliciclastic; and in rifted arc, back-arc basin, and back-arc rift tectonic settings; and are formed by metal bearing hydrothermal systems. They are defined as Au-rich if the average gold content (g/t) is greater than the combined grades of Cu, Pb, and Zn (in wt. %) and are grouped according to metallogenic association and style of mineralization. These include: A) Au-Cu deposits; B) pyritic Au deposits; and C) Au-Zn-Pb-Ag deposits. In Au-Cu deposits the mineralogical hosts to gold include native gold, Au tellurides, and auriferous pyrite; there is also a spatial correlation with advanced argillic alteration mineral assemblages (kaolinite and pyrophyllite). In pyritic Au deposits gold occurs as inclusions in arsenic-rich pyrite and arsenopyrite, and in massive pyrite zones that are low in base metal content. In Au-Zn-Pb-Ag deposits electrum, pyrite, and arsenopyrite commonly host the gold mineralization, and there is an association with feldspar alteration or gangue minerals. There are two main formational models for this deposit type: 1) syn-deformational overprinting of Au-poor base metal mineralization by metamorphic fluids; and 2) syn-volcanic mineralization by anomalously Au-rich fluids. Evidence of syn-deformational structural controls on mineralization include the location of deposits in deformed sequences proximal to regional-scale faulting and a discordant orientation of sulfide veins to regional foliation. Evidence of syn-volcanic mineralization by fluids with an anomalous chemistry include relatively un-deformed ore bodies, elevated Au concentrations over intervals greater than tens of meters, and observed cross-cutting relationships. Deposit groups classified by base metal content, Au-Cu and Au-Zn-Pb-Ag deposits, and their alteration assemblages are representative of low and high sulfidation fluids, respectively, which is analogous to hydrothermal fluid types which form epithermal type gold deposits. These formational models have been derived from studies done on, and can be applied to well-known deposits including the Archean LaRonde Penna Au-rich VMS deposit (low sulfidation), the Eskay Creek Au-Pb-Zn-Ag deposit (high-sulfidation), and the Horne Au-rich VMS deposit (Au associated with pyrite). The close distribution of Au-rich deposits to typical Au-poor base metal VMS deposits highlights the importance of understanding of formational processes to predicting the location of potential and yet undiscovered ore deposits of this type.

References

- Barrett, T. J., Cattalani, S., & MacLean, W. H., 1991. *Massive sulfide deposits of the Noranda area, Quebec. I. The Horne mine*. Canadian Journal of Earth Sciences, 28(4), 465-488.
- Dubé, B., Gosselin, P., Mercier-Langevin, P., Hannington, M.D., and Galley, A., 2007, *Gold-rich volcanogenic massive sulphide deposits*, in Goodfellow, W.D., ed., Mineral Deposits of Canada: Geological Association of Canada, Mineral Deposits Division and Geological Survey of Canada Special Publication No. 5, ISBN-13-978-1-897095-24-9, p. 75-94.
- Dubé, B., Mercier-Langevin, P., Hannington, M., Lafrance, B., Gosselin, G., & Gosselin, P., 2007. *The LaRonde Penna world-class Au-rich volcanogenic massive sulfide deposit, Abitibi, Québec: mineralogy and geochemistry of alteration and implications for genesis and exploration*. Economic Geology, 102(4), 633-666.
- Galley, A. G., M. D. Hannington, and I. R. Jonasson., 2007, *Volcanogenic massive sulphide deposits*. Mineral Deposits of Canada: A Synthesis of Major Deposit-Types, District Metallogeny, the Evolution of Geological Provinces, and Exploration Methods: Geological Association of Canada, Mineral Deposits Division, Special Publication 5 (2007): 141-161.
- Gibson, H. L., Allen, R. L., Riverin, G., & Lane, T. E., 2007. *The VMS model: Advances and application to exploration targeting*. In 5th Decennial International Conference on Mineral Exploration, Proceedings, 713-730.
- Hannington, M. D., Jonasson, I. R., Herzig, P. M., & Petersen, S., 1995. *Physical and chemical processes of seafloor mineralization at mid-ocean ridges*. Geophysical Monograph Series, 91, 115-157.
- Huston, David L. 2000. *Gold in volcanic-hosted massive sulfide deposits; distribution, genesis, and exploration*. Reviews in Economic Geology 13, : 401-426
- Marquis, P., Hubert, C., Brown, A. C., & Rigg, D. M., 1990. *Overprinting of early, redistributed Fe and Pb-Zn mineralization by late-stage Au-Ag-Cu deposition at the Dumagami mine, Bousquet district, Abitibi, Quebec*. Canadian Journal of Earth Sciences, 27(12), 1651-1671.
- Mercier-Langevin, P., B. Dube, M. D. Hannington, D. W. Davis, B. Lafrance, and G. Gosselin. 2007. *The LaRonde penna au-rich volcanogenic massive sulfide deposit, abitibi greenstone belt, quebec; part I, geology and geochronology*. Economic Geology and the Bulletin of the Society of Economic Geologists 102, (4): 585-609
- Mercier-Langevin, P., Hannington, M. D., Dubé, B., & Bécu, V., 2011. *The gold content of volcanogenic massive sulfide deposits*. Mineralium Deposita, 46(5), 509-539.
- Poulsen, K. H., & Hannington, M. D., 1996. *Volcanic-associated massive sulphide gold*. Geology of Canadian Mineral Deposit Types, Geological Survey of Canada, Geology of Canada, 8, 183-196.
- Sherlock, R. L., Roth, T., Spooner, E. T. C., & Bray, C. J., 1999. *Origin of the Eskay Creek precious metal-rich volcanogenic massive sulfide deposit; fluid inclusion and stable isotope evidence*. Economic Geology, 94(6), 803-824.
- Tourigny, G., Brown, A. C., Hubert, C., & Crepeau, R., 1989. *Synvolcanic and syntectonic gold mineralization at the Bousquet Mine, Abitibi greenstone belt, Quebec*. Economic Geology, 84(7), 1875-1890.

Convergent margin-related orogenic belts and metallogenesis

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Orogeny refers to forces and events leading to a severe structural deformation of the Earth's lithosphere due to the engagement of tectonic plates. The research of orogenic belts which are characterized by pervasive folding belts and active tectonic zones is significant since they have the potential to represent the most favorable producing areas for mineralization because of magmatic activity, faults, metamorphism pervasively took place there. In general, two genetic types including convergent-related and intraplate-related orogenic belts are recognized, in which the former orogenic belts are more essential and will be only involved in this paper. Convergent orogenic belts are further classified into continental collision-related, arc-related and accretionary orogenic belts. Continental collision occurs at convergent continental boundaries, producing mountains and suturing two continents together. Arc related orogenic belt refers to the collision between arc and continent, arc and arc, etc, and arc-continent collisional orogenic belts will be emphasized greatly here because of its remarkable role in mineralization. Accretionary orogens are the sites of long-lived convergent margin tectonics and share some similarities with collision-related orogenic belts, however, accretionary orogenic systems are represented by accreted island arc sutures and are formed in the ongoing convergent plate lasting much longer without disruption by collision. It is suggested that all these three orogenic belts would experience varied stages or geological events when considering their relationships with metallogenesis, i.e., main-collisional period, late-collisional period and post-collisional period for continental collision-related orogenic belts, constructional stage, orogenic stage and late-orogenic to post-orogenic stage for arc related orogenic belts and active subduction-related arc magmatism, superimposed rifting, inverted retro-arc pericontinental rifts, superimposed hot mantle upwellings for accretionary orogenic belts. Metallogenesis varies among different types and stages of orogenic belts is concluded and the reasons may mainly lie on diverse magma systems and fluid systems generated within related tectonic movements. In general, some magmatic hydrothermal polymetallic, porphyritic deposits and/or MVT deposits and W-Sn deposits are formed in continental collision-related orogenic belts as well as arc-related orogenic belts, however, compared with collisional orogenic belts, epithermal deposits and orogeny gold deposits are more typical in arc-related orogenic belts. Metallogenesis in accretionary orogenic belts is typically associated with gold deposits, mainly are porphyry and associated high-sulphidation epithermal Au-Cu-Ag deposits, classic low-sulphidation Au-Ag deposits, orogenic gold deposits, etc.

Keywords: orogenic belt , metallogenesis, convergence, accretionary

References

- Cawood, P.A., Kroener, A., Collins, W.J., Kusky, T.M., Mooney, W.D., Windley, B.F. 2009. Accretionary orogens through Earth history. In Cawood PA, Kroner A (eds) Earth Accretionary Systems in Space and Time. Geol Soc London, Spec Pub 318: 1–36.
- Groves, D.I., Bierlein, F.P. 2007. Geodynamic settings of mineral deposit systems. *Journal of the Geological Society, London*, 164: 19-30.
- Groves, D.I., Vielreicher, R.M., Goldfarb, R.J. & Condie, K.C. 2005b. Controls on the heterogeneous distribution of mineral deposits through time. In: McDonald, I., Noyce, A.J., Butler, I.B., Herrington, R.J. & Polya, D.A. (eds) Mineral Deposits and Earth Evolution. Geological Society, London, Special Publications, 248, 71–101.
- Hand, M., Reid, A., Jagodzinski, L. 2007. Tectonic framework and evolution of the Gawler craton, South Australia. *Econ Geol*, 92: 438–447.
- Hart, C.J.R., Mair, J.L., Goldfarb, R.J., and Groves, D.I. 2004. Source and redox controls on metallogenic variations in intrusion-related ore systems, Tombstone-Tungsten Belt, Yukon Territory, Canada: *Transactions of the Royal Society of Edinburgh, Earth Sciences*, 95, 339-356.
- Hou, Z.Q., Yang, Z.M., Qu, X.M., Meng, X.J., Li, Z.Q., Beaudoin, G., Rui, Z.Y., Gao, Y.F. 2009a. The Miocene Gangdese porphyry copper belt generated during post-collisional extension in the Tibetan orogen. *Ore Geology Reviews*, 36: 25-51.
- Hronsky, J.M.A., David, I.G., Robert R.L., Graham, C.B. 2012. A unified model for gold mineralization in accretionary orogens and implications for regional-scale exploration targeting methods. *Mineralium Deposita*, 47: 339-358.
- Jonathan, C. A., Solomon B. 2012. Accordion vs. quantum tectonics: Insights into continental growth processes from the Paleozoic of eastern Gondwana. *Gondwana Research*, 22: 674-680.
- Kerrick, R., Goldfarb, R.J. & Richards, J. 2005. Metallogenic provinces in an evolving geodynamic framework. *Economic Geology 100th Anniversary Volume*, 1097–1136.
- Leigeois, J.P. 1998. Preface—Some words on the post-collisional magmatism. *Lithos*, 45 : XV.
- Li, J.L., Sun, Shu, Hao, J., Chen, H.H., Hou, Q.L, Xiao, W.J. 1999. On the classification of collision orogenic belts. *Scientia Geologica Sinica*, 34(2):129-138.
- Saunders, J.A., Unger, D.L., Kamenov, G.D., Fayek, M., Hames, W.E., Utterback, W.C. 2008. Genesis of Middle Miocene Yellowstone hotspot-related bonanza epithermal Au-Ag deposits, Northern Great Basin, USA. *Miner Deposita*, 43:715–734.
- Sawkins, F.J. 1984. Metal deposits in relation to plate tectonics. Springer-Verlag, 2nd edition, 1-460.
- Seltman, R., Faragher, A.E. 1994. Collisional orogens and their related metallogeny—A preface. In: Seltman, R., Kampf, H., Moller, P. eds. Metallogeny of collisional orogens. Czech Geological Survey, Prague, 7-20.
- Sengor, A.M.C. 1990. Plate tectonics and orogenic research after 25 years: a Tethyan perspective. *Earth-Sci. Rev.*, 27: 1-201.
- Simmons, S.F., White, N.C. & John, D.A. 2005. Geological characteristics of epithermal precious and base-metal deposits. *Economic Geology 100th Anniversary Volume*, 485–522.
- Wang, H.Z., He, G.Q., Zhang, S.H. 2006. The geology of China and Mongolia. *Earth Science Frontiers*, 2006, 13(6): 1–13.
- Yuan, S.H., Pan, G.T., Wang, L.Q., Jiang, X.S., Yin, F.G., Zhang W.P., Zhou, J.W. 2009. Accretionary Orogenesis in the Active Continental Margins. *Earth Sciences Frontiers*, 16(3): 31-48.
- zhang, C.H. 1999. A Primary discussion on the intraplate orogenic belt. *Earth Science Frontiers*, 6(4): 295-308.

Carbonatites: A Classification and Evolutionary Review

Randy Campbell

Carbonatites were first thoroughly investigated in the late 1950's (Campbell Smith 1956; Precora 1956), and ten years later by Tuttle & Gittins (1966) who proposed some of the most problematic concepts regarding their origin. With limited advancement in the last fifty years the debate still lingers. The current scientific stalemate is in part due to the lack of extrusive carbonatites representative of their parental magmas. That being said, all carbonatites are not created equal. Currently, the IUGS classification of carbonatites allows for a wide spectrum of mineralogically and petrologically diverse rocks. This broad classification scheme requires further subdivision of carbonatites into categories that relate both their mineral chemistry and petrogenesis. This diversity has been noted by Mitchell (2005) who separates carbonatites into two groups: primary carbonatites and carbothermal residua. This review paper looks to develop a thorough understanding of their origin(s) and classification; their association with various rock types of different tectonic evolutions indicates multiple emplacement mechanisms. Using this evidence it may be possible to determine if carbonatites are sourced from a primary carbonated mantle, a result of silicate-carbonate melt immiscibility, or both. Presently there are two prime field locations where effusive carbonatites can be studied. Both Shombole (nephelinite-carbonatite) and Oldoinyo Lengai (natrocarbonatite) in East Africa indicate evidence of liquid immiscibility. This evidence is well documented and has been confirmed experimentally, others such as Harmer & Gittins (1998) would argue that ϵSr - ϵNd isotopes conclude that liquid immiscibility is not possible. It has also been shown experimentally that primary carbonatites can be generated from high magnesian melts with a total alkali content of 5-7 wt% (Harmer & Gittins, 1998). Realistically it is not possible to generate one model that satisfies the full spectrum of carbonatites. With further research it may be possible to distinguish between models, providing insight into these poorly understood magmatic/hydrothermal processes.

References

- Bailey, D. (1993). Carbonate magmas. *Journal of the Geological Society, London*, 150, 637-651.
- Bell, K., & Blenkinsop, J. (1987, February). Nd and Sr isotopic compositions of East African carbonatites: Implications for mantle heterogeneity. *Geology*, 15, 99-102.
- Bell, K., Kjarsgaard, B., & Simonetti, A. (1999). Carbonatites Into The Twenty-First Century. *Journal of Petrology*, 39, 1839-1845.
- Brooker, R. (1998). The Effect of CO₂ Saturation on Immiscibility between Silicate and Carbonate Liquids: an Experimental Study. *Journal of Petrology*, 39, 1905-1915.
- Dalton, J., & Wood, B. (1993). The compositions of primary carbonate melts and their evolution through wallrock reaction in the mantle. *Earth and Planetary Science Letters*, 119, 511-525.
- Hamilton, D., Freestone, I., Dawson, B., & Donaldson, C. (1979, May). Origin of carbonatites by liquid immiscibility. *Nature*, 279, 53-54.
- Harmer, R., & Gittins, J. (1998). The Case for Primary, Mantle-derived Carbonatite Magma. *Journal of Petrology*, 39, 1895-1903.
- Kjarsgaard, B., & Peterson, T. (1991). Nephelinite-Carbonatite Liquid Immiscibility at Shombole Volcano, East Africa: Petrographic and Experimental Evidence. *Mineralogy and Petrology*, 43, 293-314.
- Mitchell, R. (2005). Carbonatites and Carbonatites and Carbonatites. *The Canadian Mineralogist*, 43, 2049-2068.
- Nelson, D., Chivas, A., Chappell, B., & McCulloch, M. (1988). Geochemical and isotopic systematics in carbonatites and implications for the evolution of ocean-island sources. *Geochemica et Cosmochimica Acta*, 52, 1-17.
- Pecora, W. (1956, November). Carbonatites: A Review. *Review Articles in Geology*, 67, 1537-1556.
- Streckeisen, A. (1979). Classification and nomenclature of volcanic rocks, lamprophyres, carbonatites, and melilitic rocks: Recommendations and suggestions of the IUGS Subcommittee on the systematics of igneous rocks. *Geology*, 7, 331-335.
- Winter, J. (2010). *Principles of Igneous and Metamorphic Petrology* (2nd ed.). Upper Saddle River, New Jersey, United States of America: Prentice Hall.
- Woolley, A., & Kjarsgaard, B. (2008). Paragenetic Types of Carbonatites as Indicated by the Diversity and Relative Abundances of Associated Silicate Rocks: Evidence from a Global Database. *The Canadian Mineralogist*, 46, 741-752.

Bioremediation of petroleum hydrocarbon-contaminated soils: Landfarming approach

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Abstract: Bioremediation is a full-scale remediation technology which involves the use of use of micro-organisms to remove contaminants from the environment. It is broadly applicable to the remediation of petroleum hydrocarbons present in soils because hydrocarbons are biodegradable. This paper presents a review on In-situ and Ex-situ treatment processes, optimum conditions for biodegradation and effectiveness of bioremediation in clean-up of petroleum contaminated soils. Warm climates, abundant oxygen, moderate moisture content, and alkaline soil types are the favorable conditions that enhance biodegradation of Total Petroleum Hydrocarbon (TPH) present in the soil. The volume and type of contaminants are also primary factors limiting the effectiveness of bioremediation technology. Bunker C oils, a heavy fuel with complex mixtures of hydrocarbons, are recalcitrant to biodegradation and result in longer clean-up time frames. Landfarming, an onsite bioremediation technique, has been conducted on soils with moderate concentrations of hydrocarbons (~25,000mg/kg) pollutants. Substantial levels of remediation have been attained via landfarming operations within shorter time frames (4–12 week operating period) even in cold climates and remote locations. Amendments such as addition of lime to raise soil pH, bulking agents to increase aeration, and bioaugmentation significantly increase efficiency in bioremediation of petroleum contaminated soils. Enhanced bioremediation via landfarming has decreased TPH concentrations in soils by 90%, even in arctic sites with up to 4000 m³ of soils contaminated with diesel-range organics (DRO), gasoline-range organics (GRO) and BTEX compounds.

Keywords: Bioremediation, Biodegradation, Landfarming, Total Petroleum Hydrocarbons

References

- Gallego, J.R., Sierra, C., Villa, R., Pelaez, A.I., Sanchez, J., Hinrichs, K., Michaelis, W., and Rullkotter, J. 2010. Weathering processes only partially limit the potential for bioremediation of hydrocarbon-contaminated soils. *Organic Geochemistry*, **41**: 896-900.
- Kuo, Y., Wang, S., Kao, C., Chen, C., and Sung, W. 2012. Using enhanced landfarming system to remediate diesel oil-contaminated soils. *Applied Mechanics and Materials*, **121-126**: 554-558
- McCarthy, K., Walker, L., Vigoren, L., and Bartel, J. 2004. Remediation of spilled hydrocarbons by in situ landfarming at an arctic site. *Cold Regions Science and Technology*, **40**: 31-39
- Park, M. and Lee, M. 2011. Study for TPH removal efficiency of landfarming process using indigenous microorganisms to diesel contaminated site. *Mineralogical Magazine*, **75**: 1598.
- Paudyn, K., Rutter, A., Rowe, K.R., and Poland, J.S. 2008. Remediation of hydrocarbon contaminated soils in Canadian Arctic by landfarming. *Cold Regions Science and Technology*, **53**: 102-114
Proceedings - Assessment and Remediation of Contaminated Sites in Arctic and Cold Climates (ARCSACC), **3**: 257-261.
- Reisinger, H.J. 1995. Hydrocarbon bioremediation; an overview. *Bioremediation*, **3, Vol. 6**: 1-9.
- Ritter, W.F. and Scarborough, R.W. 1995. A review of bioremediation of contaminated soils and groundwater. *Journal of Environmental Science and Health, Part A: Environmental Science and Engineering*, **30**: 333-357.
- Robertson, S.J., McGill, W.B., Massicotte, H.B., and Rutherford, P.M. 2007. Petroleum hydrocarbon contamination in boreal forest soils; a mycorrhizal ecosystems perspective. *Biological Reviews (Cambridge)*, **82**: 213-240.
- Snape, I., Ferguson, S., and Revill, A. 2003. Constraints on rates of natural attenuation and in situ bioremediation of petroleum spills in antarctica.
- Lin, T., Pan, P., and Cheng, S. 2010. Ex situ bioremediation of oil-contaminated soil. *Journal of Hazardous Materials*, **176**: 27-34
- Vidali, M. 2001. Bioremediation; an overview. *Pure and Applied Chemistry*, **73**: 1163-1172.
- Westlake, D.W.S. 1999. Bioremediation, regulatory agencies, and public acceptance of this technology. *Journal of Canadian Petroleum Technology*, **38**: 48-50.
- Whyte, L.G., Goalen, B., Hawari, J., Labbe, D., Greer, C.W., and Nahir, M. 2001. Bioremediation treatability assessment of hydrocarbon-contaminated soils from eureka, nunavut. *Cold Regions Science and Technology*, **32**: 121-132.
- Zeyaulah, M., Atif, M., Islam, B., Abdelkafe, A.S., Sultan, P., ElSaady, M.A., and Ali, A. 2009. Bioremediation: A tool for environmental cleaning. *African Journal of Microbiology Research*, **3**: 310-314.

Formation of black shales: Deep versus shallow water interpretation

David B. Olotusin

Mudstones and shales are the most common sedimentary rocks. They accumulate in a variety of environments whilst comprising the bulk of recorded earth history. Previous understanding of shale formation characterised by vertical pelagic rainout from suspension is being re-evaluated. A new theory, backed by flume experimental evidence, suggests that horizontal transports are much more important. Sedimentological experimentation of flume studies has shown that mud (shale) can form deposits at flow velocities. Black shales show distinct variability in rock properties, microfabric, sub-millimetre sedimentary textures, structures, and rock properties. Petrographic evidence including thin sections reveals mud ripples, current lamination, mud intraclasts, load structures and bioturbation. These indicate the role of advective current deposit and processes in the formation of Black shales within shallow marine environments.

Black shales are organic-rich mudrock composed of silt and clay-size mineral grains. These rocks are characterized by minimum of >1% total organic carbon. Traditionally, black shales throughout the rock record were thought to have been deposited from suspension under anoxic, low-energy and quiet deep-water marine processes. These include pelagic settling, hemipelagic deposition, contourite sedimentation, turbidity current and debris flow or slides. However, recent studies have suggested that these rocks can form at any depth provided that anoxic conditions exist in water or pore fluids as well as a source of organic matter. Shale microfabric such as bedding planes, cross lamination, mud ripples, intraclasts, and bioturbation coupled with flume experimental evidence supports the idea that horizontal current transport was important. Furthermore, wave enhanced sediment flow within fair and storm weather base creates the right condition for these processes. Mechanisms of rapid settling within the storm base are responsible for remobilizing shale aggregates or sediments further seaward

This new theory has led to further research focussing on two main areas. First, at what water depth were black shales formed and secondly, the factors and processes that influenced their deposition. Laboratory investigation including petrographic evidence and thin section analysis provides direct evidence of advective current transport of mud-sized material. Clay aggregates show migrating ripples deposit sediment under higher current velocities than previously believed. Observation of current-produced particle alignment suggest that current flow over the shallow shelf was the norm rather than the exception. Also, intermittent as well as continuous current flow and reworking is indicated by sedimentary features in black shales.

These evidence confirms an interpretation of the formation of black shales within shallow marine environments. Thus, it compels a rethink or re-interpretation of existing rock and stratigraphic record with

regards to mudstones and black shales in particular. Finally, it refocuses a new understanding of black shales as a resource and how they could be better developed.

References

Macquaker, J.H. and Bohacs, K.M. 2007. Geology. on the accumulation of mud. *Science* (New York, N.Y.), **318**: 1734-1735

Macquaker, J.H.S., Bentley, S.J., and Bohacs, K.M. 2010. Wave-enhanced sediment-gravity flows and mud dispersal across continental shelves; reappraising sediment transport processes operating in ancient mudstone successions. *Geology* (Boulder), **38**: 947-950.

Plint, A.G., Macquaker, J.H.S., and Varban, B.L. 2012. Bedload transport of mud across A wide, storm-influenced ramp: Cenomanian–Turonian kaskapau formation, western canada foreland basin. *Journal of Sedimentary Research*, **82**: 801-822.

Ruppel, S.C. and Loucks, R.G. 2008. Black mudrocks; lessons and questions from the mississippian barnett shale in the southern midcontinent. *The Sedimentary Record*, **6**: 4-8.

Schieber, J., Southard, J., and Thaisen, K. 2007. Accretion of mudstone beds from migrating floccule ripples. *Science*, **318**: 1760-1763.

Schieber, J. and Yawar, Z. 2009. A new twist on mud deposition; mud ripples in experiment and rock record. *The Sedimentary Record*, **7**: 4-8.

Schieber, J., Southard, J.B., and Schimmelmann, A. 2010. Lenticular shale fabrics resulting from intermittent erosion of water-rich muds; interpreting the rock record in the light of recent flume experiments. *Journal of Sedimentary Research*, **80**: 119-128.

Schieber, J. 1994. Reflection of deep vs shallow water deposition by small scale sedimentary features and microfabrics of the chattanooga shale in tennessee. *Memoir - Canadian Society of Petroleum Geologists*, **17**: 773-784.

Stow, D.A.V., Huc, A.Y., and Bertrand, P. 2001. Depositional processes of black shales in deep water. *Marine and Petroleum Geology*, **18**: 491-498.

Tourtlot, H.A. 1979. Black shale; its deposition and diagenesis. *Clays and Clay Minerals*, **27**: 313-321.

Ver Straeten, C.A. 2012. Marcellus black shale facies; constraints and perspectives on water depth. *Abstracts with Programs - Geological Society of America*, **44**: 6.

Wignall, P.B. 2001. Shallow water, transgressive black shales. *Abstracts with Programs - Geological Society of America*, **33**: 355-356.

Wignall, P.B. 1991. Model for transgressive black shales? *Geology* (Boulder), **19**: 167-170.

Determining igneous structures and geomagnetic field history from induced and remnant magnetic fields

Colin Terry

Igneous rocks frequently have magnetic properties due to the presence of certain minerals, primarily iron and titanium oxides and iron sulfides. These properties include remnant, induced, and viscous remnant magnetization and effect the geomagnetic field of the Earth locally, producing measurable magnetic anomalies. Surveys record the magnetic field above geological structures. After correcting for diurnal, secular, anthropogenic effects, this data is used to create maps of these magnetic anomalies. Measurement of only the intensity of the magnetic field during the survey provides reasonable data for many interpretation purposes covered in this paper due to the approximation that the anomalous magnetic field is equal to the change in the local geomagnetic field so long as the International Geomagnetic Reference Field (IGRF) is much greater than the intensity of the anomaly. The magnetic field anomalies can be used to determine possible igneous structures and variations in mineral composition through a variety of processing methods. By converting the data into the frequency domain, analysis of power spectrum can provide an estimated depth to magnetic sources. The practices of reduction to pole or reduction to equator provides a means of comparing anomalies that may be due to similar geology but exist at different locations on the Earth and so appear very different in their effect on the geomagnetic field. To reduce the complexity involved in modelling a magnetic field to match surveyed results, combinations of simple structures are used to approximate dykes, vertical pipes, faults, ore bodies, and so on as prisms, tabular bodies, plates, and so on. These interpretations have significant implications in gold, diamond, and hydrocarbon exploration. Remnant magnetization records information on the direction of the geomagnetic field at the time the rock cooled past its Curie temperature. While remnant magnetization is often disregarded in favour of induced magnetic effects, it provides significant data in certain geological settings and for specific purposes. In the field of paleomagnetism these remnant magnetic vectors are to provide information on tectonic plate motion, polar reversals and true polar wander, leading to a greater understanding of the geomagnetic field of the Earth and its source, the Earth's core.

References

- Airo, M-L. Application of Aerogeophysical Data for Gold Exploration: Implications for the Central Lapland Greenstone Belt. Geological Survey of Finland (2007) Special Paper 44, 187–208
- Åm, K. The arbitrarily magnetized dyke: Interpretation by characteristics, *Geoexploration*, Volume 10, Issue 2, May 1972, Pages 63-90, ISSN 0016-7142, 10.1016/0016-7142(72)90014-2. (<http://www.sciencedirect.com/science/article/pii/0016714272900142>)
- Bean, R. J. A rapid graphical solution for the aeromagnetic anomaly of the two-dimensional tabular body. *Geophysics* 31, 963 (1966), DOI:10.1190/1.1439827
- Beckmann, G. E. J. New interpretations on palaeomagnetic data from the Nagssugtoqidian mobile belt in Greenland, *Precambrian Research*, Volume 224, January 2013, Pages 304-315, ISSN 0301-9268, 10.1016/j.precamres.2012.10.001. (<http://www.sciencedirect.com/science/article/pii/S0301926812002513>)
- Bruckshaw, J. M. and Kunaratnam, K., The interpretation of magnetic anomalies due to dykes. *Geophysical Prospecting*. July 1963
- Creveling, J. R., Mitrovica, J. X., Chan, N.-H., Latychev, K., and Matsuyama, I., Mechanisms for oscillatory true polar wander, *Nature* 491, 244–248 (08 November 2012) doi:10.1038/nature11571
- Grant, F. S., and Martin, L. Interpretation of aeromagnetic anomalies by the use of characteristic curves. *Geophysics*, February 1966, v. 31, p. 135-148, doi:10.1190/1.1439721
- Hall, D. H. A magnetic interpretation method for calculating body parameters for buried sloping steps and thick sheets, *Geoexploration*, Volume 6, Issue 4, December 1968, Pages 187-206, ISSN 0016-7142, 10.1016/0016-7142(68)90013-6. (<http://www.sciencedirect.com/science/article/pii/0016714268900136>)
- Hall, Donald H., Direction of polarization determined from magnetic anomalies. *Journal of Geophysical Research*, Volume 64, Issue 11. November 1959
- Holden E.-J., Wong, J. C., Kovesi P., Wedge, D., Dentith, M., Bagas, L. Identifying structural complexity in aeromagnetic data: An image analysis approach to greenfields gold exploration, *Ore Geology Reviews*, Volume 46, August 2012, Pages 47-59, ISSN 0169-1368, 10.1016/j.oregeorev.2011.11.002. (<http://www.sciencedirect.com/science/article/pii/S0169136811001454>)
- Moo, J. K. C., Analytical aeromagnetic interpretation the inclined prism. *Geophysical Prospecting* Volume 13. 1965. Pages 203 - 224. <http://dx.doi.org/10.1111/j.1365-2478.1965.tb01930.x>
- Reeves, C., *Aeromagnetic surveys principles, practice and interpretation*, 2005
- Steenland, N.C. Recent developments in aeromagnetic methods, *Geoexploration*, Volume 8, Issues 3–4, December 1970, Pages 185-204, ISSN 0016-7142, 10.1016/0016-7142(70)90032-3. (<http://www.sciencedirect.com/science/article/pii/0016714270900323>)
- Steinberger, B. and Torsvik, T. H., Absolute plate motions and true polar wander in the absence of hotspot tracks. *Nature* Volume 452, Issue 7187. April 2008. Pages 620 - 623

Resource Potential of the Barnett Shale, Fort Worth Basin, Texas

Tola Ogunniyi

The Newark east field of the Barnett Shale has the highest reserve of unconventional natural gas. This is as a result of temperature values greater than 450 °C, average total organic carbon (TOC) value of 3.5%, vitrinite reflectance of 1.3%, and thick accumulation of shale (> 107 meters) in this area. Shale gas is one of the major types of unconventional hydrocarbon, and shale gas plays can be found in fine grained sedimentary rocks that are rich in organic carbon. Porosity and permeability is usually low so it is almost impossible to produce gas commercially from shale without artificial stimulation or fracturing. Activities in shale gas have increased over the past two decades as there is need to look for alternative sources of hydrocarbons due to the finite nature of conventional hydrocarbons. 50% of natural gas produced in North America by 2020 will be from shale gas, and it is currently an important resource play in the United States for example where it accounted for more than 14% of gas produced as at the end 2004. The focus of this paper, which is also a world class example of a Shale gas play is the Barnett Shale in Fort Worth basin, Texas, United States. Geochemical data is important in determining the gas reserve of the Barnett shale, which has a continuous-type gas accumulation, with 2.7 trillion cubic feet (tcf) of booked reserves, and 26.22 tcf of total mean undiscovered shale gas resource. The Fort Worth basin deepens northwards and has structures that include fracturing, folds, faults (major and minor), as well as karst related collapse structures. Southern limit of the basin is defined by a dome (Llano uplift), and western boundary of the basin includes the Eastern shelf, Bend arch, and Concho platform. Muenster arches and the Red river mark the northern boundary of the basin, while the Ouachita structural front, is its margin to the east. The Barnett shale is middle - late Mississippian, and also serves as a source rock, seal and reservoir for unconventional natural gas resources. Besides, it is the largest field from which unconventional natural gas is produced in Texas. Lithofacies present in the Barnett Shale include black shale, phosphatic black shale, dolomite-rich black shale, lime grainstone and calcareous black shale. Further expansion of the area of production of the Barnett shale beyond the Newark east field have been very difficult to achieve and will require further geological, geochemical and engineering studies.

Reference List

- Bjørlykke, K. 2010. Unconventional hydrocarbons: Oil shales, heavy oil, tar sands, shale gas and gas hydrates. *In* Petroleum Geoscience: From sedimentary environments to rock physics *Edited by:* Springer, New York, pp. 459-460-465.
- Jarvie, D.M., Hill, R.J., Pollastro, R.M., Claxton, B.L., and Bowker, K.A. 2004. Evaluation of hydrocarbon generation and storage in the Barnett shale, ft. worth basin, Texas. Barnett Shale Symposium, **2**: unpaginated.
- Jarvie, D.M., Hill, R.J., and Pollastro, R.M. 2005. Assessment of the gas potential and yields from shales; the Barnett shale model. Circular - Oklahoma Geological Survey, : 37-50.
- Jarvie, D.M., Hill, R.J., Ruble, T.E., and Pollastro, R.M. 2007. Unconventional shale-gas systems; the Mississippian Barnett shale of north-central Texas as one model for thermogenic shale-gas assessment. AAPG Bulletin, **91**: 475-499.
- Loucks, R.G., Ruppel, S.C., Hill, R.J., and Jarvie, D.M. 2007. Mississippian Barnett shale; lithofacies and depositional setting of a deep-water shale-gas succession in the fort worth basin, Texas. AAPG Bulletin, **91**: 579-601.
- Montgomery, S.L., Jarvie, D.M., Bowker, K.A., and Pollastro, R.M. 2005. Mississippian Barnett shale, fort worth basin, north-central Texas; gas-shale play with multitrillion cubic foot potential. AAPG Bulletin, **89**: 155-175.
- Pollastro, R.M., Hill, R.J., Jarvie, D.M., and Henry, M.E. 2002. Assessing undiscovered resources of the Barnett-Paleozoic total petroleum system, bend arch-Fort worth basin province. American Association of Petroleum Geologists, Tulsa, OK, Tulsa, OK, United States (USA).
- Pollastro, R.M., Hill, R.J., and Jarvie, D.M. 2007. Total petroleum system assessment of undiscovered resources in the giant Barnett shale continuous (unconventional) gas accumulation, fort worth basin, Texas. AAPG Bulletin, **91**: 551-578.

The Precession of the Perihelion of Mercury

Wesley Greig

The orbit of Mercury is examined with particular attention paid to the relativistic correction to Newtonian orbital dynamics. The orbit predicted by Newtonian gravity and the effect of other planets is briefly discussed and the unexplained precession of Mercury's perihelion is investigated using the Schwarzschild solution to the Einstein equations. The geodesic equation, Killing vectors, and normalization constraints are used to derive the pertinent equations of motion. These equations are then combined to obtain an expression for the gravitational potential energy. The distinction between Newtonian gravity and that predicted by general relativity is analyzed and both are compared to observed data of Mercury's orbit. Particular attention is paid to the source of discrepancy with Newtonian gravity. The magnitude of this effect on the orbits of other planets in our Solar system is briefly investigated. The importance of relativistic corrections in gravity with respect to current research in geophysics is outlined.

References

- Böhmer, C. et al. (2010) *Classical Tests of General Relativity in Brane World models*. Classical and Quantum Gravity 27: 1 – 21.
- Clemence, G. (1947) *The Relativity Effect in Planetary Motion*. Reviews of Modern Physics 19: 361 – 364.
- Girelli, F. (2012) *Introduction to General Relativity Course Notes*. University of Waterloo.
- Iorio, L. (2011) *Gravitomagnetism and the Earth-Mercury range*. Advances in Space Research 48: 1402 – 1410.
- Rozelot, J. et al. (2009) *Probing the Solar Surface: The Oblateness and Astrophysical Consequences*. The Astrophysical Journal 703: 1791 – 1796.
- Sakhr, J. (2010) *Classical Mechanics I Course Notes*. University of Waterloo.
- Schmidt, H. (2008) *Perihelion Precession for Modified Newtonian Gravity*. Physical Review 78: 023512-1 – 023512-10.
- Standish, E. (2012) *Keplerian Elements for Approximate Positions of the Major Planets*. NASA. Retrieved November 1, 2012 from <http://ssd.jpl.nasa.gov/>
- Stewart, M. (2005) *Precession of the Perihelion of Mercury's Orbit*. American Journal of Physics 73: 730 – 734.
- Trendafilova, C. Fulling, S. (2011) *Static Solutions of Einstein's Equations with Cylindrical Symmetry*. European Journal of Physics 32: 1663 – 1677.
- Wainwright, J. (2006) *An Introduction to the Theory of Relativity*. University of Waterloo.
- Wald, R. (1984) *General Relativity*. University of Chicago Press.
- Weinberg, S. (1972) *Gravitation and Cosmology: Principles and Applications of the General Theory of Relativity*. Massachusetts Institute of Technology.
- Yan, X. et al. (2011) *Solar Oblateness and Mercury's Perihelion Precession*. Monthly Notices of the Royal Astronomical Society 415: 3335 – 3343.
- Zhao, F.; He, F. (2012) *Exact Solution to the Geodesic Motion Equation in the Braneworld Black Hole Space and the Perihelion Precession of Mercury*. International Journal of Theoretical Physics 51: 1435 – 1441.

Stromatolites: utility, application, and challenges

Mansour Al-Hashim

Stromatolites are now well-documented and known worldwide with reported occurrences from the Archean (e.g., Transvaal Supergroup, South Africa) to Recent (e.g., Shark Bay, Western Australia). They are the most well-preserved evidence of organic activities in the Precambrian. Stromatolites are widely defined as laminated biosedimentary structures formed by trapping and binding and/or precipitation of sediment particles by means of growth and metabolic activities of non-skeletal microorganisms. The main microorganisms involved in the construction of stromatolites and their structures include photosynthetic bacteria, cyanobacteria, known as blue-green algae, and filamentous, unicellular eukaryotic green algae. The morphology of stromatolites is however the product of interaction between various physical, chemical, and biological (microbial) factors that exist in the environment during their formation. This complex relationship between many different factors and the formation of stromatolites is the main source of their importance. Stromatolites are particularly useful in sedimentology, stratigraphy, paleontology, and paleoecology. They have been utilized in studying ancient depositional environments, in estimating sedimentation rates, in correlating and dating stromatolitic formations, and in regional mapping. They were also used in solving problems related to paleocurrent directions, paleolatitudes, and ancient shorelines and water depths. Understanding the microbiology and ecosystem of modern stromatolite-building biota and the processes by which modern stromatolites are being formed is critical for any sophisticated and plausible interpretation of ancient stromatolite forms. A thorough examination and description of old stromatolites is by no means less important in understanding the recent ones.

References

- Awramik, S. M., & Vanyo, J. P. (1986). Heliotropism in modern stromatolites. *Science*, 231(4743), 1279-1281.
- Des Marais, D. J. (1991). Microbial mats, stromatolites and the rise of oxygen in the precambrian atmosphere. *Global and Planetary Change*, 5(1-2), 93-96.
- Doemel, W. N., & Brock, T. D. (1974). Bacterial stromatolites; origin of laminations. *Science*, 184(4141), 1083-1085.
- Emily M, James S, Jamie S, Miriam S, Lillian, & R Pamela. (2012). Environmental controls on microbial community cycling in modern marine stromatolites. *Sedimentary Geology*, 263-264, 45-55.
- Gerdes, G., Claes, M., Dunajtschik-Piewak, K., Riege, H., Krumbein, W. E., & Reineck, H. (1993). Contribution of microbial mats to sedimentary surface structures. *Facies*, 29, 61-74.
- Hofmann, H. J., & Davidson, A. (1998). Paleoproterozoic stromatolites, hurwitz group, quartzite lake area, northwest territories, canada. *Canadian Journal of Earth Sciences = Revue Canadienne Des Sciences De La Terre*, 35(3), 280-289.
- Hofmann, H. J., Pearson, D. A. B., & Wilson, B. H. (1980). Stromatolites and fenestral fabric in early proterozoic huronian supergroup, ontario. *Canadian Journal of Earth Sciences = Revue Canadienne Des Sciences De La Terre*, 17(10), 1351-1357.
- Lanier, W. P. (1986). Approximate growth rates of early proterozoic microstromatolites as deduced by biomass productivity. *Palaaios*, 1(6), 525-542.
- Logan, B. W., Rezak, R., & Ginsburg, R. N. (1964). Classification and environmental significance of algal stromatolites. *Journal of Geology*, 72(1), 68-83.
- Logan, B. W. (1961). Cryptozoon and associate stromatolites from the recent, shark bay, western australia. *Journal of Geology*, 69(5), 517-533.
- McLoughlin, N., Wilson, L. A., & Brasier, M. D. (2008). Growth of synthetic stromatolites and wrinkle structures in the absence of microbes; implications for the early fossil record. *Geobiology*, 6(2), 95-105.
- Microbial interactions with physical sediment dynamics, and their significance for the interpretation of earth's biological history. (2008). *Geobiology*, 6(1), 93.
- Noffke, N., Knoll, A. H., & Grotzinger, J. P. (2002). Sedimentary controls on the formation and preservation of microbial mats in siliciclastic deposits; a case study from the upper neoproterozoic nama group, namibia. *Palaaios*, 17(6), 533-544.
- Schopf, J. W., Kudryavtsev, A. B., Czaja, A. D., & Tripathi, A. B. (2007). Evidence of archean life; stromatolites and microfossils. *Precambrian Research*, 158(3-4), 141-155.
- Seilacher, A. (1999). Biomat-related lifestyles in the precambrian. *Palaaios*, 14(1), 86-93.

Microfossils and the Origins of Life on Earth

Roderick Tom-Ying

Much of our understanding of life on Earth is centered on the well-studied Phanerozoic history of life where the evolution of species is well preserved in the rock strata. Unlike the Phanerozoic Eon, the basis of study for life during the Precambrian is centered on the prokaryotic microbes. Due to the nature of microbes and the poorly preserved rock record, the discussion of Precambrian fossils, as merits in inferring the origins of life are highly debated. The Apex chert formation, almost 3.5 Gya in age, found in Western Australia is a highly debated rock formation as it contains microfossils that appear biotic in origins. Held within the Precambrian chert formation is the possibility of Earth's earliest biotic microfossils. The basis of many early life hypotheses hinges on the fact that microfossils found within the rock formation appear organic in origin. Here we show that by re-examining the carbonaceous composition of the microbial filaments of said microfossils, the filament diameter, and by using Raman Spectroscopy we infer an abiotic origin. As the basis of the Early Life Hypothesis centers upon the fact that the fossils found within the Apex chert at 3.5 Gya, by re-examining the evidence we conclude that the fossils within the Apex Chert are pseudo fossils. We anticipate our results to be a starting point for a comprehensive re-examination of the Early Life Hypothesis.

References

Brasier, Martin D., Owen R. Green, John F. Lindsay, Nicola McLoughlin, Andrew Steele, and Cris Stoakes.

"Critical Testing of Earth's Oldest Putative Fossil Assemblage from the ~3.5 Ga Apex Chert, Chinaman Creek, Western Australia." *Precambrian Research* 140 (2005): 55-102. Web.

Brasier, Martin, Owen Green, John Lindsay, and Andrew Steele. "Earth's Oldest (~ 3.5 Ga) Fossils and the 'Early Eden Hypothesis': Questioning the Evidence." *Origins of Life and Evolution of the Biosphere* 34.1/2 (2004): 257-69. Print.

De Gregorio, Bradley T., Thomas G. Sharp, George J. Flynn, Sue Wirrick, and Richard L. Hervig. "Biogenic Origin for Earth's Oldest Putative Microfossils." *Geology* 37.7 (2009): 631-34. Print.

"Oldest Fossils On Earth Discovered." *ScienceDaily*. ScienceDaily, 22 Aug. 2011. Web. 01 Nov. 2012.
<<http://www.sciencedaily.com/releases/2011/08/110821205241.htm>>.

Schopf, J. William, Anatoliy B. Kudryavtsev, David G. Agresti, Thomas J. Wdowiak, and Andrew D. Czaja.
"Laser-Raman Imagery of Earth's Earliest Fossils." *Nature* 416.6876 (2002): 73-76. Print.

Schopf, J. William, Anatoliy B. Kudryavtsev, David G. Agresti, Thomas J. Wdowiak, and Andrew D. Czaja.
"Laser-Raman Spectroscopy (Communication Arising): Images of the Earth's Earliest Fossils?"
Nature 420.6915 (2002): 477. Print.

Early to Middle Ordovician Climate, Paleogeography, and Environments: Their Affects on Faunal Radiation

Kathryn Lapenskie

The Early to Middle Ordovician Earth differed greatly from the present-day planet in terms of geography, climate, environments, and marine ecosystem composition and structure. The Early to Middle Ordovician climate is characterized by greenhouse conditions, with atmospheric CO₂ concentrations up to ten times higher than modern values. High global temperatures limited or prevented the development of continental ice sheets, allowing for sea levels to attain their Phanerozoic maximum by the end of the Ordovician. The maximum continental dispersal of the last 540 Ma was achieved as four large continental landmasses, as well as microcontinents and island arcs, were widely distributed throughout the southern hemisphere. The northern hemisphere was unoccupied by continents and covered by the vast Panthalassic Ocean.

Extensive epicontinental seas developed in the southern hemisphere due to globally high sea levels. Wide temperate and tropical marine belts developed as a result of a greenhouse climate. Sluggish ocean circulation limited the upwelling of deeper, nutrient rich waters, causing superoligotrophic conditions. Hardgrounds and flat sea beds became common throughout the shallow cratonic seas. Island arcs and ocean terranes provided fauna platforms on which to migrate and radiate on.

The Great Ordovician Biodiversification Event was a significant faunal radiation, occurring in a 25 million year interval during the Early to Middle Ordovician. Several climatic, environmental, and tectonic factors led to the development of this event. Geographic isolation of organisms, due to fragmented continents and intense tectonic activity, drove speciation and diversity. The extensive tropical, shallow epicontinental seas created by the greenhouse conditions and high sea levels provided habitats for organisms to thrive in. Increased primary productivity in cratonic seas enabled a subsequent diversity of primary producers.

Organisms with mineralized skeletons became highly diversified during this radiation, including brachiopods, bryozoans, cephalopods, conodonts, solitary and colonial corals, echinoderms, graptolites, ostracodes, sponges, and trilobites. Reef composition changed from microbially- to metazoan-dominated framework builders. New niches were exploited as organisms occupied different tiers above and below the sediment-water interface. Planktonic animals expanded their environments to inhabit greater ranges of the water column. The poorly organized ecosystems of the Cambrian Period, dominated by epifaunal animals, were replaced by complex, predictable food webs. The new families, genera, and species arising out of the Great Ordovician Biodiversification Event compose the Paleozoic Evolutionary Fauna, which dominated marine communities until the end of the Permian Period.

References

- Algeo, T.J. and Sessler, K.B. 1995. The Paleozoic world: continental flooding, hypsometry, and sealevel. *American Journal of Science*, **295**: 787-822.
- Barnes, C.R. 2004. Ordovician oceans and climate. *In* The Great Ordovician Biodiversification Event. *Edited by* B.D. Webby, F. Paris, M.L. Droser, and I.G.Percival. Columbia University Press, New York, N.Y. pp. 72-76.
- Cocks, L.R.M. and Torsvik, T.H. 2004. Major terranes in the Ordovician. *In* The Great Ordovician Biodiversification Event. *Edited by* B.D. Webby, F. Paris, M.L. Droser, and I.G.Percival. Columbia University Press, New York, N.Y. pp. 61-67.
- Harper, D.A.T. 2006. The Ordovician biodiversification: Setting an agenda for marine life. *Palaeogeography, Palaeoclimatology, Palaeoecology*, **232**: 148-166.
- Herrmann, A.D., Patzykowski, M.E., and Pollard, D. 2004a. The impact of paleogeography, ρCO_2 , poleward ocean heat transport and sea level change on global cooling during the Late Ordovician. *Palaeogeography, Palaeoclimatology, Palaeoecology*, **206**: 59-74.
- Herrmann, A.D., Haupt, B.J., Patzkowski, M.E., Seidov, D., and Slingerland, R.L. 2004b. Response of Late Ordovician paleoceanography to changes in sea level, continental drift, and atmospheric ρCO_2 : Potential causes for long-term cooling and glaciations. *Palaeogeography, Palaeoclimatology, Palaeoecology*, **210**:385-401.
- Miller, K.G., Kominz, M.A., Browning, J.V., Wright, J.D., Mountain, G.S., Katz, M.E., Sugarman, P.J., Cramer, B.S., Christie-Blick, N., and Pekar, S.F. 2005. The Phanerozoic record of global sea-level change. *Science*, **310**: 1293-1298.
- Munnecke, A., Calner, M., Harper, D.A.T., and Servais, T. 2010. Ordovician and Silurian sea-water chemistry, sea level, and climate: A synopsis. *Palaeogeography, Palaeoclimatology, Palaeoecology*, **296**: 389-413.
- Servais, T., Owen, A.W., Harper, D.A.T., Kröger, B., and Munnecke, A. 2010. The Great Ordovician Biodiversification Event (GOBE): The palaeoecological dimension. *Palaeogeography, Palaeoclimatology, Palaeoecology*, **294**: 99-119.

**Recognition of tsunami deposits within the sedimentary record:
Attempts at establishing a tsunami facies**

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Tsunami deposit identification within the sedimentary record is difficult due to characteristics such as preservation potential, grain size, and depositional environment, causing the appearance of the deposit to vary. To date, tsunami deposits have yet to be identified based on sedimentary criteria alone, requiring dating of sediments immediately overlying each deposit followed by a comparison to a record of historical tsunami occurrence. The establishment of a sedimentary facies unique to tsunamis has implications for further deposit identification and, by extension, risk assessment.

A review of scientific research on tsunami deposits from different parts of the world, including Japan, Australia, and Portugal, have provided a baseline of common sedimentary characteristics, as observed through trenching and coring techniques. Deposits are commonly divided into four sub-units, each of which is interpreted to correspond to a different phase of the tsunami; the initial, smaller waves of the tsunami depositing the lowest most unit, followed by the large, powerful waves, the waning energy waves, and finally the post-tsunami fall-out. Sedimentary characteristics of the deposits have included landward thinning and fining, rip-up clasts, cross-bedding, grading, and boulders, to name a few. This paper provides an overview of features common to tsunami deposits as well as limitations associated with their discovery and interpretation.

References

- Fujiwara, O. 2008. Bedforms and Sedimentary Structures Characterizing Tsunami Deposits. *In* Tsunamiites: Features and Implications. *Edited by* T. Shiki, Y. Tsuji, T. Yamazaki, and K. Minoura. Elsevier, Oxford, U.K. pp. 51–62.
- Keating, B.H., Wanink, M., and Helsley, C.E. 2008. Introduction to a Tsunami-Deposits Database. *In* Tsunamiites: Features and Implications. *Edited by* T. Shiki, Y. Tsuji, T. Yamazaki, and K. Minoura. Elsevier, Oxford, U.K. pp. 359–381.
- Kortekaas, S., and Dawson, A.G. 2007. Distinguishing tsunami and storm deposits: An example from Martinhal, SW Portugal. *Sedimentary Geology* **200**: 208–221.
- Mamo, B., Strotz, L., and Dominey-Howes, D. 2009. Tsunami sediments and their foraminiferal assemblages. *Earth-Science Reviews* **96**: 263–278.
- Peters, R., Jaffe, B., and Gelfenbaum, G. 2007. Distribution and sedimentary characteristics of tsunami deposits along the Cascadia margin of western North America. *Sedimentary Geology* **200**: 372–386.
- Phantuwongraj, S., and Choowong, M. 2012. Tsunamis versus storm deposits from Thailand. *Natural Hazards* **63**: 31–50.
- Ramírez-Herrera, M-T., Lagos, M., Hutchinson, I., Kostoglodov, V., Machain, M.L., Caballero, M., Goguitchaichvili, A., Aguilar, B., Chagué-Goff, C., Goff, J., Ruiz-Fernández, A-C., Ortiz, M., Nava, H., Bautista, F., Lopez, G.I., and Quintana, P. 2012. Extreme wave deposits on the Pacific coast of Mexico: Tsunamis or storms? – A multi-proxy approach. *Geomorphology* **139–140**: 360–371.
- Shanmugam, G. 2012. Process-sedimentological challenges in distinguishing paleo-tsunami deposits. *Natural Hazards* **63**: 5–30.

A Comparison of two Projects for the Prevention of High Water in the City of Venice, Italy, as a Result of Land Subsidence and Climate-Induced Sea Level Rise

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The frequency and intensity of flooding in the city of Venice, Italy, has increased over the past 50 years. This phenomenon is governed by a combination of geological, historical and climatological factors. This review will aim to summarize the geomorphological evolution of the Venetian Lagoon and of the land subsidence phenomenon, meanwhile providing a background of the historical evolution of the area. This will include the main anthropogenic causes of flooding in Venice and the recent attempts and future options to save Venice from rising water levels. In particular, we will examine the two most discussed projects: the Experimental Electromechanical Module (MoSE project), and the Anthropogenic Uplift of Venice City . The MoSE project was approved by the Italian government in 2001 and is currently under construction. It consists of movable barriers installed in correspondence with the three lagoonal inlets and is designed to block seawater inflow into the lagoonal basin during exceptional tidal events that cause flooding in Venice. With the current sea level, the activation of the movable barriers is expected to occur with a frequency of 2-3 times per year but may increase as a result of climate-induced sea level rise. This will have a negative impact on harbour activities in Venice as maritime traffic will be limited at the inlets. As a result, MoSE project might become obsolete in less than 100 years. A complementary project has been proposed: the Anthropogenic Uplift of Venice. It consists of 12 vertical wells strategically located within the lagoon, that inject seawater into the 600-800 deep aquifer. The numerical model described here, predicts an uplift of between 11 and 40 cm over a 10 year period. Preliminary results shows that the anthropogenic Uplift of Venice might be a promising complementary action to MoSE barriers as it has the potential to reduce the frequency of the closure of the inlets by prolonging the operational life of MoSE.

References

- Brambati, A., Carbognin, L., Quaia, T., Teatini, P., and Tosi, L. 2003. The Lagoon of Venice: geological setting, evolution and land subsidence. *Episodes-News magazine of the International Union of Geological Sciences*, **26**(3): 264-268.
- Canu, D.M., Umgiesser, G., Solidoro, C. 2001. Short-term simulations under winter conditions in the lagoon of Venice: a contribution to the environmental impact assessment of temporary closure of the inlets. *Ecological Modeling* **138**: 215–230.
- Carbognin, L., and Tosi, L. 2002. Interaction Between Climate Changes, Eustacy and Land Subsidence in the North Adriatic Region, Italy. *Marine Ecology*, **23**: 38-50.

- Carbognin, L. and Tosi, L. 2003. Il progetto ISES per l' analisi dei processi di intrusione salina e subsidenza nei territori meridionali delle province di Padova e Venezia, ISMAR Institute of Marine Sciences, Venice, Italy.
- Carbognin, L., Marabini, F. and Tosi, L. 1995. Land subsidence and degradation of the Venice littoral zone, Italy. *Journal of Marine Systems* **51**: 345–353.
- Carbognin, L., Teatini, P., Tomasin, A., and Tosi, L. 2010. Global change and relative sea level rise at Venice: what impact in term of flooding. *Climate Dynamics*, **35**(6): 1039-1047.
- Castelletto, N., Ferronato, M., Gambolati, G., Putti, M., and Teatini, P. 2008. Can Venice be raised by pumping water underground? A pilot project to help decide. *Water Resources Research*, **44**(1): 1-16.
- Comerlati, A., Ferronato, M., Gambolati, G., Putti, M., and Teatini, P. 2004. Saving Venice by Seawater. *Journal of Geophysical Research* **109**: 1-14.
- Fontini, F., Umgiesser, G., and Vergano, L. 2010. The role of ambiguity in the evaluation of the net benefits of the MOSE system in the Venice lagoon, *Ecological Economics* **69**: 1964–1972.
- Jamiolkowski, M., G. Ricceri, and Simonini, P. 2009. Safeguarding Venice from high tides: site characterization & geotechnical problems. *In Proceedings of the 17th International Conference on Soil Mechanics and Geotechnical Engineering*, Alexandria, Egypt.
- Pirazzoli, P. A. 2002. Did the Italian Government approve an obsolete project to save Venice?. *Eos, Transactions American Geophysical Union*, **83**(20): 217-223.
- Tomasin, A., and Frassetto, R. 2001. Cyclogenesis and Forecast of Water Dynamic Elevations in Venice. *In Marine Forecasting: Predictability and Modeling in Ocean Hydrodynamics*. Edited by Nihoul J.C. Elsevier Scientific Publishing Company, New York, pp. 427-438.
- Umgiesser, G., Canu, D.M., Cucco, A., and Solidoro, C. 2004. A finite element model for the Venice Lagoon. Development, set up, calibration and validation, *Journal of Marine Systems* **51**: 123–145.
- Umgiesser, G., and Matticchio, B. 2006. Simulating the mobile barrier (MOSE) operation in the Venice Lagoon, Italy: global sea level rise and its implication for navigation. *Ocean Dynamics*, **56**(3): 320-332.
- Teatini, P., Ferronato, M., Gambolati, G., Baù, D., and Putti, M. 2010. Stochastic Analysis of the Venice uplift due to Seawater Injection into Deep Aquifers. *In XVIII International Conference on Water Resources*, Barcelona, Spain.

Effects of Climate-Induced Temperature and Water Table Changes on Carbon-Dynamics of Northern Peatlands

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Peatlands account for only 3% of the Earth's land area; however they are an important net carbon sink, with northern peatlands storing up to 30% of the world's carbon stocks. Climate change is expected to have a significant impact on the structure and physicochemical characteristics of peatlands, especially those at high latitudes. Of the projected consequences, warmer temperatures and water table drawdown are of particular concern, each having direct and indirect effects on the carbon-dynamics of peat soils. This review will aim to provide a general summary of how changing environmental conditions will impact various aspects of the carbon cycle in northern peatlands, specifically soil and gaseous carbon. Increased temperature and lower water table levels have been found to enhance decomposition rates of organic matter in soils, which would effectively increase carbon dioxide emissions to the atmosphere. Warmer temperatures will provide favourable temperatures for methanogenesis, however a lowering of the water table below 10-20 cm and change in plant community structure could decrease the amount of methane that is emitted to the atmosphere. The drivers of dissolved organic carbon concentrations in peatlands are not yet known, which creates uncertainty in the estimation of how dissolved organic carbon responds to climatic disruptions. Drought has been found to decrease dissolved organic carbon concentrations owing to increased mineralization rates, but greater temperatures have been noted to have both positive and negative effects on concentrations. This overview suggests that climate-induced warming and drying of northern peatlands will increase carbon dioxide and decrease methane effluxes from soils, however no clear consensus for soil carbon components has been established.

References

- Blodau, C. 2002. Carbon cycling in peatlands: A review of processes and controls [online]. *Environmental Reviews*, **10**: 111-134. doi=10.1139%2Fa02-004
- Bridgham, S. D., Johnston, C. A., Pastor, J., and Updegraff, K. 1995. Potential feedbacks of northern wetlands on climate change [online]. *BioScience*, 262-274. Available from <http://www.jstor.org/stable/1312419> [accessed 8 October 2012].
- Davidson, E. A., and Janssens, I. A. 2006. Temperature sensitivity of soil carbon decomposition and feedbacks to climate change [online]. *Nature*, **440**: 165-173. doi:10.1038/nature04514
- Gorham, E. 1991. Northern peatlands: role in the carbon cycle and probable responses to climatic warming [online]. *Ecological applications*, **1**: 182-195. Available from <http://www.jstor.org/stable/1941811> [accessed 30 October 2012].
- Knorr, K. H., and Blodau, C. 2009. Impact of experimental drought and rewetting on redox transformations and methanogenesis in mesocosms of a northern fen soil [online]. *Soil Biology and Biochemistry*, **41**: 1187-1198. doi:10.1016/j.soilbio.2009.02.030
- Laudon, H., Buttle, J., Carey, S. K., McDonnell, J., McGuire, K., Seibert, J., Shanley, J., Soulsby, C. and Tetzlaff, D. 2012. Cross-regional prediction of long-term trajectory of stream water DOC response to climate change [online]. *Geophysical Research Letters*, **39**: L18404. doi:10.1029/2012GL053033
- Limpens, J., Berendse, F., Blodau, C., Canadell, J. G., Freeman, C., Holden, J., Roulet, N., Rydin, H. and Schaepman-Strub, G. 2008. Peatlands and the carbon cycle: from local processes to global implications - a synthesis [online]. *Biogeosciences Discussions*, **5**: 1379-1419. Available from <http://hal.archives-ouvertes.fr/hal-00297992/> [accessed 31 October 2012].
- Preston, M. D., Eimers, M. C., and Watmough, S. A. 2011. Effect of moisture and temperature variation on DOC release from a peatland: Conflicting results from laboratory, field and historical data analysis [online]. *Science of the Total Environment*, **409**: 1235-1242. doi:10.1016/j.scitotenv.2010.12.027
- Riley, J.L. 2011. *Wetlands of the Ontario Hudson Bay Lowlands: A Regional Overview*. Nature Conservancy of Canada, Toronto, ON.
- Trettin, C. C., Laiho, R., Minkinen, K., and Laine, J. 2006. Influence of climate change factors on carbon dynamics in northern forested peatlands [online]. *Canadian journal of soil science*, **86**: 269-280. doi: 10.4141/S05-089
- Ye, R., Jin, Q., Bohannon, B., Keller, J. K., McAllister, S. A., and Bridgham, S. D. 2012. pH controls over anaerobic carbon mineralization, the efficiency of methane production, and methanogenic pathways in peatlands across an ombrotrophic-minerotrophic gradient [online]. *Soil Biology and Biochemistry*, **54**: 36-47.

Millennial Climate Cycles in the Holocene

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Future climate projections have been constrained to what can be concluded of the mechanisms that drove climate change in the past. As climate change becomes a more global public concern, the need for a complete understanding of natural climate cycles is critical before the recent anthropogenic impact can be interpreted relative to the natural variability.

Holocene climate, the most recent interglacial period, was thought to be relatively stable until paleo records uncovered a common millennial periodicity. Glacial and interglacial cycles driven by Milankovitch orbital forcing operate on a scale of $> 20,000$ years. However, these cycles can be punctuated by shorter cycles of lower amplitude such as 11-year Schwabe cycles, 85-year Gleissburg cycles, and 207-year deVries cycles. Recently, a new millennial climate cycle of ~ 1400 years has been evidenced by various paleo records and suggest a much larger global impact mechanism.

Carbon-14 and Beryllium-10, both cosmogenically created proxies for total solar irradiance (TSI), have been correlated to these millennial climate cycles. This evidence suggests that solar forcing may contribute to a new time scale of climate variability in the natural environment. Based on isotopes, pollen, and *foraminifera* paleo data, it has been hypothesized that solar forcing influences the atmosphere enough to change the sea surface temperature, the North Atlantic Oscillation (NAO), and even the thermohaline circulation of the ocean. The resulting change in North Atlantic Deep Water (NADW) formation amplifies this solar forcing to a global scale and may have implications for our current climate-warming event.

Other mechanisms have been presented in the literature including volcanism and glacial influence. Thus, more research will need to be done in order to conclusively determine the main driving force of this millennial oscillation. Linking these mechanisms of climate change is crucial to understanding Earth's past climate regimes and making predictions for Earth's future climate.

References

- Arzel, O., de Verdiere, A.C., and England, M.H. 2010. The role of oceanic heat transport and wind stress forcing in abrupt millennial-scale climate transitions. *Journal of Climate*, **23**: 2233-2256.
- Bond, G., Kromer, B., Beer, J., Muscheler, R., Evans, M., Showers, W., Hoffmann, S., Lotti-Bond, R., Hajdas, I., and Bonani, G. 2001. Persistent solar influence on north atlantic climate during the holocene. *Science*, **294**: 2130-2136.
- Bond, G., Showers, W., Cheseby, M., Lotti, R., Almasi, P., deMenocal, P., Priore, P., Cullen, H., Hajdas, I., and Bonani, G. 1997. A pervasive millennial-scale cycle in north atlantic holocene and glacial climates. *Science*, **278**: 1257-1266.
- Flint, R.F. 1971. *Glacial and quaternary geology*.
- Grove, J.M. 1979. The glacial history of the holocene. *Progress in Physical Geography*, **3**: 1-54.
- Hamanaka, N., Kan, H., Yokoyama, Y., Okamoto, T., Nakashima, Y., and Kawana, T. 2012. Disturbances with hiatuses in high-latitude coral reef growth during the holocene: Correlation with millennial-scale global climate change. *Global and Planetary Change*, **80-81**: 21-35.

- Ivanochko, T., Ganeshram, R., Brummer, G., Ganssen, G., Jung, S., Moreton, S., and Kroon, D. 2005. Variations in tropical convection as an amplifier of global climate change at the millennial scale. *Earth and Planetary Science Lett*, **235**: 302-314.
- Jouzel, J., Masson-Delmotte, V., Cattani, O., Dreyfus, G., Falourd, S., Hoffmann, G., Minster, B., Nouet, J., Barnola, J.M., Chappellaz, J., Fischer, H., Gallet, J.C., Johnsen, S., Leuenberger, M., Loulergue, L., Luethi, D., Oerter, H., Parrenin, F., Raisbeck, G., Raynaud, D., Schilt, A., Schwander, J., Selmo, E., Souchez, R., Spahni, R., Stauffer, B., Steffensen, J.P., Stenni, B., Stocker, T.F., Tison, J.L., Werner, M., and Wolff, E.W. 2007. Orbital and millennial antarctic climate variability over the past 800,000 years. *Science*, **317**: 793-796.
- Kloosterboer-van Hove, M., Steenbrink, J., Visscher, H., and Brinkhuis, H. 2006. Millennial-scale climatic cycles in the early pliocene pollen record of ptolemais, northern greece. *Palaeogeography Palaeoclimatology Palaeoecology*, **229**: 321-334.
- Kohfeld, K. and Harrison, S. 2000. How well can we simulate past climates? evaluating the models using global palaeoenvironmental datasets. *Quaternary Science Reviews*, **19**: 321-346.
- Li, Y., Yu, Z., and Kodama, K.P. 2007. Sensitive moisture response to holocene millennial-scale climate variations in the mid-atlantic region, USA. *Holocene*, **17**: 3-8.
- Lu Yingxia, Li Baosheng, Wen Xiaohao, Qiu Shifan, Wang Fengnian, Niu Dongfeng, and Li Zhiwen. 2010. Millennial-centennial scales climate changes of holocene indicated by magnetic susceptibility of high-resolution section in salawusu river valley, china. *Chinese Geographical Science*, **20**: 243-251.
- Matthews, J.A. and Dresser, P.Q. 2008. Holocene glacier variation chronology of the smorstabtindan massif, jotunheimen, southern norway, and the recognition of century- to millennial-scale european neoglacial events. *Holocene*, **18**: 181-201.
- McManus, J.F., Bond, G.C., Broecker, W.S., Johnsen, S., Labeyrie, L., and Higgins, S. 1994. High-resolution climate records from the north atlantic during the last interglacial. *Nature*, **371**: 326-329.
- McManus, J.F., Oppo, D.W., and Cullen, J.L. 1999. A 0.5-million-year record of millennial-scale climate variability in the north atlantic. *Science*, **283**: 971-975.
- Oppo, D.W., Keigwin, L.D., McManus, J.F., and Cullen, J.L. 2001. Persistent suborbital climate variability in marine isotope stage 5 and termination II. *Paleoceanography*, **16**: 280-292.
- Poore, R., Dowsett, H., Verardo, S., and Quinn, T. 2003. Millennial- to century-scale variability in gulf of mexico holocene climate records. *Paleoceanography*, **18**: 1048.
- Raymo, M.E., Ganley, K., Carter, S., Oppo, D.W., and McManus, J. 1998. Millennial-scale climate instability during the early pleistocene epoch. *Nature*, **392**: 699-702.
- Schulz, M., Paul, A., and Timmermann, A. 2004. Glacial-interglacial contrast in climate variability at centennial-to-millennial timescales: Observations and conceptual model. *Quaternary Science Reviews*, **23**: 2219-2230.
- Shimada, C., Ikehara, K., Tanimura, Y., and Hasegawa, S. 2004. Millennial-scale variability of holocene hydrography in the southwestern okhotsk sea: Diatom evidence. *Holocene*, **14**: 641-650.
- Steenbrink, J., Kloosterboer-van Hove, M., and Hilgen, F. 2003. Millennial-scale climate variations recorded in early pliocene colour reflectance time series from the lacustrine ptolemais basin (NW greece). *Global and Planetary Change*, **36**: 47-75.
- Stuiver, M., Grootes, P., and Braziunas, T. 1995. The GISP2 delta O-18 climate record of the past 16,500 years and the role of the sun, ocean, and volcanoes. *Quaternary Research*, **44**: 341-354.
- Thompson, S., Govindasamy, B., Mirin, A., Caldeira, K., Delire, C., Milovich, J., Wickett, M., and Erickson, D. 2004. Quantifying the effects of CO₂-fertilized vegetation on future global climate & carbon dynamics. *Geophys. Res. Lett.* **31**: L23211
- Weber, M.E., Tougiannidis, N., Kleineder, M., Bertram, N., Ricken, W., Rolf, C., Reinsch, T., and Antoniadis, P. 2010. Lacustrine sediments document millennial-scale climate variability in northern greece prior to the onset of the northern hemisphere glaciation. *Palaeogeography Palaeoclimatology Palaeoecology*, **291**: 360-370.
- Yu, Z., Campbell, I., Campbell, C., Vitt, D., Bond, G., and Apps, M. 2003. Carbon sequestration in western canadian peat highly sensitive to holocene wet-dry climate cycles at millennial timescales. *Holocene*, **13**: 801-808.
- Zhao, C., Yu, Z., Zhao, Y., Ito, E., Kodama, K.P., and Chen, F. 2010. Holocene millennial-scale climate variations documented by multiple lake-level proxies in sediment cores from hurleg lake, northwest china. *J. of Paleolimnology*, **44**: 995-1008.

Strength of the Continental Lithosphere

Mengmeng Qu

The strength of the continental lithosphere, or maximum stress it can support before failing, is crucial in geodynamics. Its spatial and temporal variations can help us understand the Earth's deformation processes including rifting, mountain building, sedimentary basin development, seismicity and volcanism.

In the 1970s, Goetze and Evans firstly introduced the yield stress envelope (YSE), a vertical profile predicting the maximum differential stress supported by rock as a function of depth for the oceanic lithosphere. This concept works well for the oceanic lithosphere, because it can explain the response of observed age and temperature dependence of plate to surface and subsurface loads. But when it comes to the continental lithosphere, problems appear. Compared with oceanic lithosphere, continental lithosphere is in a much more complicated context. It has a thicker crust and a longer deformation history, and bears the modification by surficial process (e.g., erosion, sedimentation and orogenesis). In the 1980s, based on the study of the distribution of focal depths for earthquakes, Chen and Molnar stated the classical view on the strength of the continental lithosphere: the continental lithosphere generally consisted of a weak lower crust sandwiched between a relatively strong upper crust and uppermost mantle. This is known as “jelly sandwich”.

However, at the beginning of the 21st century, after the reassessment of earthquake depth distributions and gravity anomalies, Jackson and Maggi found that there was little support in earthquake focal depth distributions, for the uppermost mantle was significantly stronger than the lower crust in continental regions. Therefore they proposed an opposite view, “crème brûlée”, suggesting that the strength of the continental lithosphere resided in the crust, and that the upper mantle beneath the continents was relatively weak. To analyze which idea is more applicable, in accordance with them, Burov used dynamic numerical models to test the stability and structural styles. The results turned out to be compatible with the view that the lithospheric mantle was strong (“jelly sandwich”) and in this way, the continental lithosphere could support geological loads and stress for long periods of time. Therefore, they concluded that “jelly sandwich” was more widely applicable.

In the paper, I review the two opinions about the strength of the continental lithosphere and focus on recent researches: 1) Jackson and Maggi's study on the focal depth distribution of earthquakes and the association of gravity anomalies with topography; 2) Burov's dynamic numerical models. Then I analyze problems in these researches and general difficulties in studies of the strength of the continental lithosphere.

In the end, I give my own understanding of the strength of the continental lithosphere and perspectives on future studies.

References

- Burov E.B. . 2010. *The equivalent elastic thickness (T_e), seismicity and the long-term rheology of continental lithosphere: Time to burn-out “crème brûlée”?* Insights from large-scale geodynamic modeling, *Tectonophysics*, 484: 4–26
- Burov E.B. , Michel Diament. 1995. *The effective elastic thickness (T_e) of continental lithosphere: What does it really mean?*, *Journal of Geophysical Research*, v. 100, no. B3: 3905–3927
- Burov E.B., Watts A.B. . 2006. *The long-term strength of continental lithosphere: “jelly sandwich” or “crème brûlée”?*, *GSA Today*, v. 16, no. 1: 4–10
- Chen Wang-Ping, Hung Shu-Huei , Tai-Lin Tseng, Michael Brudzinski, Zhaohui Yang, Robert L. Nowack. 2012. *Rheology of the continental lithosphere: Progress and new perspectives*, *Gondwana Research*, 21: 4–18
- Chen Wang-Ping, Peter Molnar. 1983. *Focal depths of intracontinental and intraplate earthquakes and their implications for the thermal and mechanical properties of the lithosphere*, *Journal of Geophysical Research*, v. 88, no. B5: 4183–4214
- Jackson James. 2002. *Strength of the continental lithosphere: Time to abandon the jelly sandwich?*, *GSA Today* (Spet. 2002): 4–9
- Karato Shun-ichiro . 2010. *Rheology of the deep upper mantle and its implications for the preservation of the continental roots: A review*, *Tectonophysics*, 481: 82–98
- Karato Shun-ichiro. 2011. *Some Issues on the Strength of the Lithosphere*, *Journal of Earth Science*, v. 22, no. 2: 131–136
- Kohlstedt D. L., Brian Evans, Mackwell S.J. . 1995. *Strength of the lithosphere by laboratory experiments*, *Journal of Geophysical Research*, v.100, no. B9: 17587–17602
- Maggi A., Jackson J.A. , D. McKenzie, K. Priestley. 2000. *Earthquake focal depths, effective elastic thickness, and the strength of the continental lithosphere*, *Geology*, v. 28; no. 6: 495–498
- Naliboff J. B., C. Lithgow-Bertelloni, L. J. Ruff , N. de Koker. 2012. *The effects of lithospheric thickness and density structure on Earth’s stress field*, *Geophys. J. Int.* 188: 1–17
- Poliakov, A.N.B., Cundall, P., Podladchikov, Y., Laykhovsky, V., 1993. *An explicit inertial method for the simulation of visco-elastic flow: an evaluation of elastic effects on diapiric flow in two- or three-layers models*. In: Stone, D.B., Runcorn, S.K. (Eds.), *Flow and Creep in the Solar System: Observations, Modelling and Theory*, *Dynamic Modeling and Flow in the Earth and Planets Series*: 175–195
- Tesauro Magdala, Mikhail K. Kaban, Sierd A.P.L. Cloetingh. 2012. *Global strength and elastic thickness of the lithosphere*, *Global and Planetary Change*, 90–91: 51–57
- Tesauro Magdala, Pascal Audet, Mikhail K. Kaban, Roland Bürgmann, Sierd Cloetingh. 2012. *The effective elastic thickness of the continental lithosphere: Comparison between rheological and inverse approaches*, *Geochem. Geophys. Geosyst.*, 13, Q09001
- Toussaint, G., Burov, E., and Jolivet, L.. 2004a. *Continental plate collision: Unstable vs. stable slab dynamics*, *Geology*, v. 32: 33–36
- Toussaint, G., Burov, E., Avouac, J.-P., 2004b. *Tectonic evolution of a continental collision zone: a thermo-mechanical numerical model*. *Tectonics* 23, TC6003.

Shale gas in Canada: geological controls and current challenges

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Abstract: Natural gas accounts for a significant part of the energy consumption around the world. With the increasing concerns about unconventional natural gas, the exploration and exploitation of shale gas has developed dramatically in recent years, and it result in the increase of natural gas reserves. The term “Shale Gas” refers to unconventional, continuous-type, self-sourced resources contained in fine grained (ranging from clay to very fine sandstone), organic-rich, low permeability reservoirs in which thermogenic or biogenic methane is stored as free gas in the matrix or fracture porosity, or as adsorbed/dissolved gas on the organics and/or clays. These are self-enclosed petroleum systems, characterized by inefficient “dysfunctional” expulsion and migration, where source, reservoir and trap are all present in the same thick shaly succession. The most prospective shale gas targets will be thick, widespread, gas-saturated, fine grained, organicrich units. There are three main geological controls for shale gas plays: tectono-stratigraphic position, organic matter content and reservoir characteristics. Thickness and distribution area is the key conditions which ensure there is enough storage space and organic matter. There is positive correlation between the organic matter content and the methane capacity of shale. The features of pore and fracture determine it whether we can get access to the economic flow and how we design the project for hydraulic fracturing. Besides, challenges for the development of shale gas in Canada are analyzed in the end. The intrinsic characteristics of the shale gas and the fundamental controls on its productivity need to be well understood. Currently there are few production wells in Canada with a history of well performance that can be used to extrapolate recoverable potential. The assessment methodology for continuous resource must be flexible enough to adapt to situations ranging from little or no well data to thousands of production wells. Modern geological information, beginning with maps, is needed for evaluation of shale gas targets.

Keywords: Shale gas; Geological controls; Canada; Challenges

References

- American Association of Petroleum Geologists. 2011. Unconventional Energy Resources: 2011 Review. Natural Resources Research, **20(4)**: 297.
- Ball, C. 2005b. Shale Silence Deafening: potential of shale gas drives behind the scenes plans; Oil week, **August**: 23-25.

- Bustin, A.M.M., Bustin, R.M. 2012. Importance of rock properties on the producibility of gas shales. *International Journal of Coal Geology* **103**: 132–147.
- Bustin, R.M. 2005. Gas shale tapped for big pay; *AAPG Explorer*, **February**.
- Bustin, R.M. 2006. Geology Report: where are the high-potential regions expected to be in Canada and the U.S.? *Second Annual Capturing Opportunities in Canadian Shale Gas*, Calgary.
- Chalmers, G.R.L., Bustin, R.M. 2006. The organic matter distribution and methane capacity of the Lower Cretaceous strata of Northeastern British Columbia, Canada. *International J. Coal Geology* **70**: 223–239.
- Chalmers G.R.L., Ross J.K., Bustin, R.M. 2012. Geological controls on matrix permeability of Devonian Gas Shales in the Horn River and Liard basins, northeastern British Columbia, Canada. *International Journal of Coal Geology* (In press).
- Curtis, J.B. 2002. Fractured shale-gas systems; *American Association of Petroleum Geologists Bull.* **86**: 1921-1938.
- Faraj, B. 2005. Critical Elements of Shale Gas Exploration; *Seventh Unconventional Gas Conference*, Calgary.
- Faraj, B. 2006. Shale Gas activity update for Canada and the U.S.: who is doing what...and where? *Second Annual Capturing Opportunities in Shale Gas Conference*, Calgary.
- Frantz, J. 2006. Exploration innovation: getting the most out of exploration technology to know what you have and develop winning shale gas play; *Second Annual Capturing Opportunities in Canadian Shale Gas Conference*, Calgary.
- Hamblin, A.P.. 2006. The “Shale Gas” concept in Canada: a preliminary inventory of possibilities. *GEOLOGICAL SURVEY OF CANADA, OPEN FILE 5384*.
- Lavoie, D., Chen, Z., Pinet, N., Lyster, S.. 2012. A review of November 24-25, 2011 shale gas workshop. *GEOLOGICAL SURVEY OF CANADA, OPEN FILE 7088*.
- Schein, G. 2006. Well stimulation success stories: Barnett shale completions; *Second Annual Capturing Opportunities in Canadian Shale Gas Conference*, Calgary.
- Schurr, G.W., Ridgley, J.L. 2002. Unconventional shallow biogenic gas systems; *American Association of Petroleum Geologists Bulletin* **86**: 1939-1969.
- Shirley, K. 2004. Unorthodox plays can muddy roles; *Am. Association of Petroleum Geologists Explorer*, **July**.
- Waters, G. 2006. Considerations for shale gas completions; *Second Annual Capturing Opportunities in Canadian Shale Gas Conference*, Alta.

Groundwater remediation using zero-valent iron as a reactive medium in permeable reactive barriers

Alana Crump

Conventional methods of groundwater remediation involve pumping water to the surface where subsequent treatment occurs, followed by the release of this treated water back into an aquifer. However, these pump-and-treat systems are energy and maintenance intensive, invoking the need to replace traditional technologies. Permeable reactive barriers (PRBs) have recently emerged as a viable method for passive remediation of contaminated groundwater. These barriers are installed in the path of a contaminated groundwater plume and contain reactive materials that promote various geochemical reactions. PRBs therefore transform contaminants into innocuous components as groundwater flows through a subsurface diaphragm under the natural hydraulic gradient. Adsorption or redox reactions followed by the precipitation of sparingly-soluble compounds are examples of processes occurring in PRBs, however the ability to predict long-term performance of PRBs in different hydrogeochemical environments is difficult.

The use of Zero-Valent Iron as a reactive medium is of particular interest for the treatment of numerous contaminants. This material is the most-widely used medium for permeable reactive barriers, and laboratory and field experiments have proven its effectiveness. Meanwhile, several experiments have demonstrated that the formation of secondary precipitates often reduces the permeability of PRBs over time. Although current studies advocate the use of this method as a replacement for groundwater remediation, additional research must be conducted regarding the long-term effects of these precipitates on porosity and hydraulic conductivity. This review therefore outlines PRB technology, provides a list of contaminants that are treatable using Zero-Valent Iron, and summarizes problematic aspects of this method that have been determined in laboratory and field research.

References

- Benner, S.G., Blowes, D.W., Gould, W.D., Herbert Jr, R.B., and Ptacek, C.J. 1999. Geochemistry of a permeable reactive barrier for metals and acid mine drainage. *Environmental Science and Technology*. **33**: 2793-2799.
- Blowes, D.W., Ptacek, C.J., and Jambor, J.L. 1997. *In-situ* remediation of Cr(VI)-contaminated groundwater using permeable reactive walls: laboratory studies. *Environmental Science and Technology*. **31**(12): 3348-3357.
- Blowes, D.W., Ptacek, C.J., Benner, S.G., McRae, C.W.T., Bennett, T.A., and Puls, R.W. 2000. Treatment of inorganic contaminants using permeable reactive barriers. *Journal of Contaminant Hydrology*. **45**: 123-137.
- Calabro, P.S., Moraci, N., and Suraci, P. 2012. Estimate of the optimum weight ratio in Zero-Valent Iron/Pumice granular mixtures used in permeable reactive barriers for the remediation of nickel contaminated groundwater. *Journal of Hazardous Materials*. **207-208**: 111-116.
- Guo Q., and Blowes, D.W. 2009. Biogeochemistry of two types of permeable reactive barriers, organic carbon, and iron-bearing organic carbon for mine drainage treatment: column experiments. *Journal of Contaminant Hydrology*. **107**: 128-139.
- Jeen, S-W, Gillham, R.W., and Blowes, D.W. 2006. Effects of carbonate precipitates on long-term performance of granular iron for reductive dechlorination of TCE. *Environmental Science and Technology*. **40**: 6432-6437.
- Lindsay, M.B.J., Ptacek, C.J., Blowes, D.W., and Gould, W.D. 2008. Zero-valent iron and organic carbon mixtures for remediation of acid mine-drainage: Batch experiments. *Applied Geochemistry*. **23**: 2214-2225.
- Ludwig, R.D. Smyth, D.J.A, Blowes, D.W., Spink, L.E., Wilkin, R.T., Jewett, D.G., and Weisener, C.J. 2009. Treatment of arsenic, heavy metals, and acidity using a mixed ZVI-compost PRB. *Environmental Science and Technology*. **43**: 1970-1976.
- Miao, Z., Brusseau, M.K., Carroll, K.C., Carreon-Diazconti, C., and Johnson, B. 2012. Sulfate reduction in groundwater: characterization and applications for remediation. *Environmental and Geochemical Health*. **34**: 539-550.
- Richardson, J.P., and Nicklow, J.W. 2002. *In situ* permeable reactive barriers for groundwater contamination. *Soil and Sediment Contamination*. **11**(2): 241-268.
- Wilkin, R.T., Su, C., Ford, R.G., and Paul, C.J. 2005. Chromium-removal processes during groundwater remediation by a zerovalent iron permeable reactive barrier. *Environmental Science and Technology*. **39**: 4599-4605.

Mining salt, brine and clay: A review of lithium and boron evaporite deposits

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Hydrogen fuel cells and battery powered electric vehicles are at the cusp of focus for electrical power storage as their need for portable power and energy is directly relative to their success. Elements such as lithium and now boron have been targeted as candidates to solving chemical-electrical storage problems. Lithium and boron are extracted from continental nonmarine evaporite deposits that comprise 82.7% and 90% of the world supply respectively (Dundee Capital Markets, 2009) (Smith & Medrano, 1996) and are extracted principally as salts, brines and clays (Warren, 1999).

The following paper is a review of lithium and boron evaporite deposits focusing on their origins, hydrogeochemical controls, extraction and economics. Deposits are segregated based on lithium-boron bearing mineralogy or brine chemistry. Depositional facies, stratigraphy and diagenesis are overviewed and illustrate a complex hydrogeological system dependent on water-rock interactions.

Economic parameters of deposit types efficiency, sustainability and economic potential are defined. Findings converge that lithium brines are the most economical source of lithium carbonate still but offer significant setbacks to product to market timing because of solar dependence for concentrating brines and the variability of brine in aquifers. In contrast, boron production is strongly concentrated in Turkey and the U.S.A with 72% of total reserves located in Turkey (Kar, et al. 2006) and offers a stable supply platform. Conversely, economic potentials of evaporite deposits are solely dependent on demand and emerging technologies. Boron is viewed as having a higher economic potential because of emerging uses of borohydrides in hydrogen fuel cells.

References

- Dundee Capital Markets. 2009. Lithium - Hype or Substance? A look at Lithium Demand and Supply. Dundee Securities Corporation.
- Kar, Y., Sen, N., Demirbas, A. 2006. Boron Minerals in Turkey, Their Application Areas and Importance for the Country's Economy. *Minerals and Energy*, **20**, 2-10.
- Smith, G. I., and Medrano, M. D. 1996. Continental Borate Deposits of Cenozoic Age. *Mineralogical Society of America*, **33**, 263-284.
- Warren, J. 1999. *Evaporites, Their Evolution and Economics*. Blackwell Scienc Ltd., Malden, MA.

Self-organized criticality: What can it tell us about natural hazards?

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Self-organized criticality is a phenomenon that describes a system evolves to a critical state spontaneously rather than by external fine tuning of a parameter as in phase transition. A system described by self-organized criticality exhibits power-law behaviour and fractal size distribution of events. The concept of self-organized criticality evolved from the study of three major types of models: the sand-pile model, describing the evolution of a sandpile by consistently adding sands; the forest fire model, which involves the dynamics of burning trees and the avalanches in the system are forest fires; and the slider-block model, which is a simple analogue for the behaviour of faults in the Earth's crust through blocks motion driven by the friction force of a plate. Self-organized criticality has been successful in describing complexity, together with chaos and fractals. It has been proven to be a new approach to study a wide range of complex systems from large scale natural phenomena to human social behaviour, e.g. landslides, earthquakes, forest fires, brain activity, stock markets, epidemics.

Some natural hazards such as landslides, forest fires and earthquakes are characterized by unpredictable events, or avalanches, as well as a power-law scaling of frequency-size distribution. For example, the Gutenberg-Richter law, which describes the size distribution of earthquakes, exhibits characteristics of self-organized criticality. Simple self-organized criticality models have been proven to display strong descriptive power and can be directly applied to natural systems. The aim of this paper is to introduce the framework of self-organized criticality and review the applications of models to natural hazards. The sand-pile models have been applied to landslides and rockfalls, the forest fire models to forest fires and wild fires, the slider-block models to earthquakes. In addition, potential self-organized criticality models for volcanic eruption are discussed. The model behaviour yields to a good analogue of the actual observations, the estimates of the size and frequency of possible events can be drawn. Moreover, acquiring more insight into the mechanism of natural hazards through investigating the model behaviour could further foster constructing hazard and risk assessment systems.

Self-organized criticality is still at an early stage of development and not a well-defined concept. It is controversial in the sense that the process of self-organizing into the critical state and the mathematical proof of the power-law behaviour is not quite clear. Nonetheless, self-organized criticality serves as an approach to an in-depth understanding of the dynamics of dissipative non-equilibrium systems. Questions remain as to whether or not it is really a 'universal' behaviour and truly captures the essence of the phenomenon.

References

- Bak, P., Tang, C., Wiesenfeld, K. 1987. *Self-organized Criticality: An Explanation of 1/f Noise*, Phys. Rev. Lett., **59**: 4.
- Bak, P., Tang, C., Wiesenfeld, K. 1988. *Self-organized criticality*, Phys. Rev. A, **38**: 1.
- Bak, P., Tang, C. 1989. *Earthquakes as a Self-Organized Critical Phenomenon*, J.G.R. **94**: 15635-15637.
- Dhar, D. 2006. *Theoretical studies of self-organized criticality*, Physica A, **369**: 2970.
- Diodatia, P., Bak, P., Marchesoni, F. 2000. *Acoustic emission at the Stromboli volcano scaling laws and seismic activity*, Earth and Planetary Science Letters, **182**: 253-258.
- Drossel, B., Schwabl, F., 1992. *Self-organized critical forest-fire model*, Phys. Rev. Lett., **69**: 11.
- Frette, V., Christensen, K., Malthe-Sorensen, A., Feder, J. 1996. *Avalanche dynamics in a pile of rice*, Nature, **379**: 6560.
- Hergarten, S., Krenn, R. 2011. *Synchronization and desynchronization in the Olami-Feder-Christensen earthquake model and potential implications for real seismicity*, Nonlin. Processes Geophys., **18**: 635642.
- Hergarten, S. 2012. *Topography-based modeling of large rockfalls and application to hazard assessment*, Geophysical Research Letters, **39**: L13402.
- Jensen, H.J. 1998. *Self-Organized Criticality: Emergent Complex Behavior in Physical and Biological Systems*, Cambridge University Press.
- Krenn, R., Hergarten, S. 2009. *Cellular automaton modelling of lightning-induced and man made forest fires*. Nat. Hazards Earth Syst. Sci., **9**: 17431748.
- Lee, W. H. K. 2009. *Introduction to Complexity in Earthquakes, Tsunamis, and Volcanoes, and Forecast*, in Encyclopedia of Complexity and Systems Science, Springer, pp 68-78.
- Malamud, B.D., Millington, J. D. A., Perry, G. L. W., Turcotte, D. L. 2005. *Characterizing Wildfire Regimes in the United States*, in Proceedings of the National Academy of Sciences of the United States of America, Vol. 102, No. 13, pp. 4694-4699.
- Sachs, M. K., Yoder, M. R., Turcotte, D. L., Rundle, J. B., Malamud, B. D. 2012. *Black swans, power laws, and dragon-kings: Earthquakes, volcanic eruptions, landslides, wildfires, floods, and SOC models*, Eur. Phys. J. Special Topics, **205**: 167-182.
- Song, W., Wang, J., Satoh, K., Fan, W. 2006. *Three types of power-law distribution of forest fires in Japan*, Ecological Modeling, **196**: 527532.
- Turcotte, D. L. 2001. *Self-organized criticality: Does it have anything to do with criticality and is it useful?* Nonlinear Processes in Geophysics, **8**: 193196.
- Van Den Eeckhaut, M., Poesen, J., Govers, G., Verstraeten, G., Demoulin, A. 2007. *Characteristics of the size distribution of recent and historical landslides in a populated hilly region*, EPSL. **256**: 588-603.

Basin Controls on the occurrence of Reservoir Intervals in the Cardium Formation, AB

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Mudstone and sandstone are the major rocks that can be found in Cardium formation where outcrops the Rockies foothills and lies beneath the Alberta plain. Conglomerate fractions could also be found in this formation not a lot but important. Sediments were accumulated in the area of muddy and sandy marine environment. The process of autocyclic and allocyclic controlled depositional basin. Cardium sediments are contained by a large bow- shaped basin, which follows an elongate compound arc that trends northwest.

The tectonic system as the Tintina-Northern Rocky Mountain Trench (TT-NRMT) fault system transformed from right-lateral strike- slip to compressional deformation basement subsidence increased in the southern Canadian Rockies. As orogenesis occurred large amounts of sediment were delivered to the subsiding basin contributing to overall subsidence. High-frequency fluctuations in sea level due to eustatic and tectonic controls resulted in the complex depositional patterns of the Cardium formation. .

The Cardium formation is a prolific hydrocarbon deposit producing both conventional and unconventional resources. Basin controls on reservoir units are mainly tectonic and also eustatic. Orogenesis caused subsidence in the Alberta Basin creating a large sediment source and accommodation space to deposit ‘the prograding clastic wedge of the Cardium formation. Relative changes in sea level caused an erosional unconformity and back stepping, which combined with gravel input and wave reworking created some of the best reservoirs in the Cardium formation. Conglomerate and high quality sand reservoirs were originally conventional targets. Presently low permeability ‘fringe’ or ‘halo’ deposits around high quality reservoirs are popular unconventional horizontal fracturing targets.

Keywords: basin; subsidence; eustatic; prograding clastic; conventional; unconventional

References

- Arnott, R.W.C. 2003. The role of fluid- and sediment-gravity flow processes during deposition of deltaic conglomerates (Cardium Formation, Upper Cretaceous), west- central Alberta. *Bulletin of Canadian Petroleum Geology*, v. 51, p. 426–436.
- Dashtgard, S.E., Buschkuehle, M.B.E, Fairgrieve, B. and Berhane, H. 2008. Geological characterization and potential for carbon dioxide (CO₂) enhanced oil recovery in the Cardium Formation, central Pembina Field, Alberta. *Bulletin of Canadian Petroleum Geology*, v. 56, no. 2, p. 147-164
- Hall, R.L., Krause, F.F., Joiner, S.D. and Deutsch, K.B. 1994. Biostratigraphic evaluation of a sequence stratigraphic surface: the Cardinal/Leyland unconformity (“E5/T5 surface”) in the Cardium Formation (Upper Cretaceous; upper Turonian-lower Coniacian) at Seebe, Alberta. *Bulletin of Canadian Petroleum Geology*, v. 42, no. 3, p. 296-311
- Hart, B.S. and Plint, A.G. 2003. Stratigraphy and sedimentology of shoreface and fluvial conglomerates: insights from the Cardium Formation in NW Alberta and adjacent British Columbia. *Bulletin of Canadian Petroleum Geology*, v. 51, no. 4, p. 437-464
- Krause, F.F., Deutsch, K.B., Joiner, S.D., Barclay, J.E., Hall, R.L. and Hills, L.V. 2010. Cretaceous Cardium Formation of the Western Canada Sedimentary Basin. In: *Geological Atlas of the Western Canada Sedimentary Basin*, G.D. Mossop. and I. Shetsen (compilers), Canadian Society of Petroleum Geologists and Alberta Research Council, Special Report 4
- Leckie, D.A., Bhattacharya, J.P., Bloch, J., Gilboy, C.F. and Norris, B. 2010. Cretaceous Colorado/Alberta Group of the Western Canada Sedimentary Basin. In: *Geological Atlas of the Western Canada Sedimentary Basin*, G.D. Mossop. and I. Shetsen (compilers), Canadian Society of Petroleum Geologists and Alberta Research Council, Special Report 4
- Leshchyshyn, T.T. and Pierre-Gilles, B. 2010. Multistage Fractured Cardium Oil: Studying Geology and All of the Production and Stimulation History Thus Far. Canadian Society for Unconventional Gas, Society of Petroleum Engineers, CSUG/SPE 137678
- National Energy Board, 2011. Tight Oil Developments in the Western Canada Sedimentary Basin. Energy Briefing Note, p. 38
- Plint, A.G., Walker, R.G. and Duke, W.L. 1988. An outcrop to subsurface correlation of the Cardium Formation in Alberta. In: *Sequences, Stratigraphy, Sedimentology: Surface and Subsurface*. D.P. James and D.A. Leckie (eds.). Canadian Society of Petroleum Geologists, Memoir 15, p. 167-184
- Price, R.A. 2010. Cordilleran Tectonics and the Evolution of the Western Canada Sedimentary Basin. In: *Geological Atlas of the Western Canada Sedimentary Basin*, G.D. Mossop. and I. Shetsen (compilers), Canadian Society of Petroleum Geologists and Alberta Research Council, Special Report 4
- Walker, R.G. 1983a. Cardium Formation 1. “Cardium a turbidity current deposit” (Beach 1955): A brief history of ideas. *Bulletin of Canadian Petroleum Geology*, v. 31, p. 205-212
- Walker, R.G. and Eyles, C.H. 1990. Topography and Significance of a Basinwide Sequence-Bounding Erosion Surface in the Cretaceous Cardium Formation, Alberta, Canada. *Journal of Sedimentary Petrology*, v. 61, no. 4, p. 473-496

Seismic risk in Canada

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The seismic hazard exists almost anywhere in the world, and Canada is no exception. Every day there are about 3-4 earthquakes in this country, and in spite of the fact that most of them can only be detected by sensitive equipment, a few times a year, Canadians really feel the movement of the earth. The urgency of this issue highlights by the fact that a strong earthquake near large cities in Canada can produce damage in the billions of dollars and lead to the deaths of thousands of people. Most seismically hazardous territory of Canada is Western Canada, particularly British Columbia. Seismic activity in this region caused earthquakes with magnitudes greater than 8 Richter. Also in eastern Canada there is a risk of seismic hazard. The evidence is devastating earthquakes up to magnitude 7 Richter, occurred near the St. Lawrence and Ottawa rivers and earthquakes near North Bay, Ontario, in January 2000 and the earthquake in Georgian Bay in October 2005. In central Canada there is no significant risk of earthquakes. That is why in this paper will be considered a historical aspect of seismic and seismic risk in Western and Eastern Canada.

The analysis of seismic risk is determined as the consequences and the likelihood of hazardous events that can happen in that period of time, which is one of the most important tasks of management and is used to analyze the economic, social and environmental consequences of hazardous events. In the context of this paper will discuss the main theoretical approaches to the analysis and evaluation of seismic risk, and methods to reduce it. Determination of seismic risk in Canada plays an important role in the upcoming events, which can greatly help in the assessment of risk and response planning, mitigate losses and tragedies associated with these extremes, and reduce the effects of seismic hazard on the Canadian citizens and infrastructure of Canadian cities.

References

- Adams, John; Rogers, G; McCormack, D; Cassidy, J; Lamontagne, M. 2001. An advanced national earthquake monitoring system for Canada's cities at risk. International Association of Seismology and Physics of the Earth's Interior: International Association of Geomagnetism and Aeronomy.
- Atkinson, Gail. 1982. Application of current seismic risk analysis methods to critical facilities. Washington, DC, United States (USA): U. S. Natl. Sci. Found.
- Aylsworth, J M. 2012. New Canadian teaching resources from Natural Resources Canada for earthquakes, landslides and tsunamis. Open-File Report: Geological Survey of Canada: 7.
- Emmi, Philip C; Horton, Carl A. 1996. Seismic risk assessment, accuracy requirements, and GIS-based sensitivity analysis. CO, United States (USA): GIS World Books.
- Kluegel, Jens-Uwe; Attinger, Richard; Panza, G F. January 2011. Scenario-based seismic risk analysis; an engineering approach to the development of source and site-specific ground motion time histories in areas of low seismicity. *Pure and Applied Geophysics*, **168**: 55-67.
- Kovacs, Paul. 2004. Awareness of seismic risk in Eastern Canada. *Seismological Research Letters*, **75**: 452.
- Leboeuf, D; Nollett, M J; Jolicoeur, L; Achim, C. February 2007. Microzonation and vulnerability assessment for seismic risk management in Quebec City. *Seismological Research Letters*, **78**: 158.
- Mazzotti, S; Leonard, L J; Cassidy, J F; Rogers, G C; Halchuk, S. 2011. Seismic hazard in Western Canada from GPS strain rates versus earthquake catalog. *Journal of Geophysical Research* **116**: Citation B12310.
- Nuta, Elena; Christopoulos, Constantin; Packer, Jeffrey A. March 2011. Methodology for seismic risk assessment for tubular steel wind turbine towers; application to Canadian seismic environment. *Canadian Journal of Civil Engineering*, **3**: 293-304.
- Onur, Tuna; Ventura, Carlos E; Finn, W D Liam. Apr. 2005. Regional seismic risk in British Columbia - damage and loss distribution in Victoria and Vancouver. *Canadian Journal of Civil Engineering*, **2**: 361-371.
- Patchett, Annette; Robinson, David; Dhu, Trevor; Sanabria, Augusto. 2004. Incorporating uncertainty in probabilistic seismic risk analyses. Australia (AUS): Australian Earthquake Engineering Society.
- Ploeger, S K; Atkinson, G M; Samson, C. April 2010. Applying the HAZUS-MH software tool to assess seismic risk in downtown Ottawa, Canada. *Natural Hazards*, **1**: 1-20.
- Spence, R J; Coburn, A W; Sakai, S; Pomonis, A. Barking. 1991. A parameterless scale of seismic intensity for use in seismic risk analysis and vulnerability assessment. Essex, United Kingdom.
- Stevens, Anne E. 1988. Earthquake hazard and risk in Canada. Dordrecht, Netherlands : D. Reidel Publ. Co.
- Tiampo Kristy. 2007. Predicting Canada's Next Earthquake. Available from <http://www.canadianunderwriter.ca/news/predicting-canada-s-next-earthquake/1000211164/>.
- Varley, Paul M. 1996. Seismic risk assessment and analysis. London, United Kingdom: Thomas Telford.

Uses and Challenges in Real-Time seismological data applications

Behzad Hassani

Real-time seismological data applications refer to the applications in which seismic data are received and analyzed quickly after a significant seismic event, so that the results can be effectively used for post-earthquake emergency response and early warning. In general, real-time applications are based on two distinct procedures. In first one, the rapid (3-5 minutes) generation of maps of instrumental ground-motion (acceleration, velocity, and spectral response) and shaking intensity accomplished thorough near real-time seismographic data acquisition combined with developed relationships between recorded ground-motion parameters and expected shaking intensity values. These automatic maps which are triggered by any significant earthquake can provide a rapid depiction of the extent of potentially damaging area following an earthquake and can be used for emergency response, loss estimation and for public information through the media. As the technology of seismic instrumentation, telemetry, computers, and data storage facility advances, the real-time seismology for rapid post-earthquake notification is essentially established.

The second procedure which is called Earthquake Early Warning (EEW) is based on the idea that seismic waves radiate at a lower speed than electromagnetic waves that are used to transmit possible warnings in case of strong events. This leading time can be implemented to reduce likely damages that might be caused by the later arriving seismic waves. Research for EEW is currently in progress and two methods are widely used: (1) regional warning and (2) on-site warning methods. In regional warning method, the traditional seismological techniques are used to locate an earthquake, determine the magnitude and estimate the ground motion at other sites. In on-site warning method, the beginning of the ground motion (P wave) observed at a site is used to predict the following ground motion (S wave) at the same site and no attempt is necessarily made to locate an earthquake and estimate the magnitude. The first approach is more reliable, but it takes a longer time and cannot be used for the sites at short epicentral distances. However, the second approach is less accurate, but it is very fast and can provide useful early warning to sites even at very short distances where an early warning is most vital. To benefit from the advantages of both methods some integrated approaches have been proposed to not only use the accuracy of the first approach but also benefit from the fast responses of the second approach. The uses of EEW can be considered both at personal and institutional levels. For instance, Personal protective measures can be undertaken at home and in the workplace include getting under desks and moving away from dangerous chemicals and machinery. At institutional level, protective measures can be exploited at mass-transportation systems that can use a few seconds to slow and stop trains, terminate airplane landings, and prevent additional cars from entering the freeway.

References

- Allen, R.M., Kanamori, H. 2003. *The potential for earthquake early warning in southern California*, Science, **300**: 786–9.
- Allen, R.M., Gasparini, P., Kamigaichi, O., Böse, M. 2009. *The status of earthquake early warning around the world: an introductory overview*, Seismological Research Letter, **80(5)**: 682-693.
- Böse, M. 2006. *Earthquake Early Warning for Istanbul using Artificial Neural Networks*, PhD thesis, Karlsruhe University, Germany.
- Böse, M., Ionescu, C., Wenzel, F. 2007. *Earthquake early warning for Bucharest, Romania: Novel and revisited scaling relations*, Geophysical Research Letters, **34**: 1-6.
- Böse, M., Wenzel, F., Erdik, M. 2008. *PreSEIS: A neural network based approach to earthquake early warning for finite faults*, Bulletin of the Seismological Society of America, **98**:366–382.
- Hsiao, N.C., Wu, Y.M., Shin, T.C., Zhao, L., Teng T.L. 2009. *Development of earthquake early warning system in Taiwan*, Geophysical Research Letters, **36**, L00B02.
- Jalpa, D. P., Atkinson, M.G. 2012. *Scenario ShakeMaps for Ottawa, Canada* . Bulletin of the Seismological Society of America, 102(2): 650–660.
- Kanamori, H. 2005. *Real-time seismology and earthquake damage mitigation*, Annual Review of Earth and Planetary Sciences, **33**: 195–214.
- Kanamori, H., 2008, *Earthquake physics and real-time seismology*, Nature, **451**: 271-3.
- Kanamori, H.; Hauksson, E., Heaton, T. 1997. *Real-time seismology and earthquake hazard mitigation*, Nature, **390**: 461-4.
- Moratto, G.C., Suhadolc, P. 2009. *Real-Time Generation of ShakeMaps in the Southeastern Alps*, Bulletin of the Seismological Society of America, **99(4)**: 2489–2501.
- Wu, Y.M., Kanamori, H. 2005. *Experiment on an onsite early warning method for the Taiwan early warning system*, Bulletin of the Seismological Society of America, **95**: 347–353.
- Wu, Y. M., Kanamori, H. 2008. *Development of an earthquake early warning system using real-time strong motion signals*, Sensors., **8**: 1–9.
- Wald, D.J., Quitoriano, V., Heaton, T.H., Kanamori, H., Scrivner, C.W., Worden, C.B. 1999. *TriNet “ShakeMaps”: rapid generation of peak ground motion and intensity maps for earthquakes in southern California*, Earthquake Spectra, **15**:537–55.
- Wald, D., Wald, L., Worden, B., Goltz, J. 2007. *USGS ShakeMap—A tool for earthquake response*, U.S. Geological Survey, Fact Sheet 087-03-508, 4 pp.

Detachment faulting and its implications for the mineralization of Oceanic Core Complexes

Jon Hey

Scientific interest in oceanic core complexes (OCC) has greatly increased since the initial mapping expedition to the Atlantis Massif in 1996. Ongoing research has focused on mantle structure, marine magnetic anomalies, OCC formation and their relation to detachment faults. OCCs form on the inside corner of ridge transform-fault intersections along slow spreading to ultra-slow spreading ridges. Current models support long-lived, large-scale detachment faults, creating axial asymmetry. Reduced magma supply and the exposure of lower crustal and mantle rocks suggest extension involving predominantly tectonic instead of magmatic processes. The hanging wall of these faults are typically metamorphosed ultramafic schists, with a serpentinized peridotite in the footwall, intruded by discrete gabbroic bodies. The fault comprises a network of smaller anastomosing fault zones, generally dipping at $\leq 20^\circ$. The fault gouge is typically 1 – >200 m thick. Recent studies have shown that hydrothermal mineralization has been occurring at temperatures of 300-400°C up to 12 km off-axis along the Mid-Atlantic Ridge. New oxygen and strontium isotopic evidence has shown that active oceanic detachment faults can focus large volumes of hydrothermal fluids, and are the primary conduits for these fluids and slow spreading ridges. Geophysical studies of the Trans-Atlantic Geotraverse (TAG) have shown that the detachment fault dips at roughly 20° towards the ridge axis until a depth of 1 km, at which point it plunges at 70° to a depth of >7 km. There is also no geophysical evidence for any crustal melt reservoirs at shallower depths to provide heat to the fluids. The evolution of detachment fault mineralization can be broken down into three stages: 1) early, intense hydrothermal circulation, driven by hot gabbroic intrusions into serpentinized ultramafic footwall rocks. TAG-type vents occur here during final discharge with a preference for Fe-Cu-Zn-Si mineralization in basalt; 2) fluid flow through the mature detachment has fluids interacting with both gabbroic intrusions and serpentinized peridotite, discharging through ultramafic rich footwall rocks at $\sim 370^\circ\text{C}$ in Rainbow-type vents, showing high temperature Cu-Zn-Fe-Co-Au-(Ni) sulfide mineralization in ultramafics; and 3) low temperature circulation in cooled peridotites distal to the ridge axis generates low-temperature Si-(Zn-Cu) and Ca-Mg deposits in Lost City-type venting. The Cu-Zn-Co-Au deposits are more common in OCCs than in ophiolites, suggesting that ultramafic hosted volcanogenic massive sulfide deposits on slow spreading ridges fail to accrete during obduction and are thus a type of mineralization specific to the marine environment. Based on the extensional nature of OCC formation, it may be possible to find obducted complexes in failed rift zones accreted to continents.

References

- Blackman, D. K., Cann, J. R., & Janssen, B. (1998). Origin of extensional core complexes: evidence from the Mid-Atlantic Ridge at Atlantis Fracture Zone. *Journal of Geophysical Research*, 103(B9), 21315–21333.
- Canales, J. P., Sohn, R. A., & deMartin, B. J. (2007). Crustal structure of the Trans-Atlantic Geotraverse (TAG) segment (Mid-Atlantic Ridge, 26°10'N): Implications for the nature of hydrothermal circulation and detachment faulting at slow spreading ridges. *Geochemistry Geophysics Geosystems*, 8(8), Q08004. doi:10.1029/2007GC001629
- Cann, J. R., Blackman, D. K., Smith, D. K., & McAllister, E. (1997). Corrugated slip surfaces formed at ridge-transform intersections on the Mid-Atlantic Ridge. *Nature*, 385, 329–332.
- deMartin, B. J., Sohn, R. A., Canales, J. P., & Humphris, S. E. (2007). Kinematics and geometry of active detachment faulting beneath the Trans-Atlantic Geotraverse (TAG) hydrothermal field on the Mid-Atlantic Ridge. *Geology*, 35(8), 711–714.
- Fouquet, Y., Cambon, P., Etoubleau, J., Charlou, J. L., Ondréas, H., Barriga, F. J. A. S., Cherkashov, G., et al. (2010). Geodiversity of Hydrothermal Processes Along the Mid-Atlantic Ridge and Ultramafic-Hosted Mineralization: A New Type of Oceanic Cu-Zn-Co-Au Volcanogenic Massive Sulphide Deposit. In P. A. Rona, C. W. Devey, J. Dymont, & B. J. Murton (Eds.), *Diversity of Hydrothermal Systems on Slow Spreading Ocean Ridges* (Vol. 188, pp. 321–367). Washington, D. C.: American Geophysical Union. doi:10.1029/2008GM000746
- Hedenquist, J. W., & Lowenstern, J. B. (1994). The role of magmas in the formation of hydrothermal ore deposits. *Nature*, 370, 519–527.
- Ildefonse, B., Blackman, D. K., John, B. E., & Ohara, Y. (2007). Oceanic core complexes and crustal accretion at slow-spreading ridges. *Geology*, 35(7), 623–626.
- James, R. H., & Elderfield, H. (1996). Chemistry of ore-forming fluids and mineral formation rates in an active hydrothermal sulfide deposit on the Mid-Atlantic Ridge. *Geology*, 24(12), 1147–1150.
- Karson, J. A. (1999). Geological investigation of a lineated massif at the Kane Transform Fault: Implications for oceanic core complexes. *Philosophical Transactions of the Royal Society of London A*, 357, 713–740.
- Kelley, D. S., Karson, J. A., Früh-Green, G. L., Yoerger, D., Shank, T. M., Butterfield, D., Hayes, J. M., et al. (2005). A Serpentinite-Hosted Ecosystem: The Lost City Hydrothermal Field. *Science*, 307(5714), 1428–1434. doi:10.1126/science.1102556
- Lowell, R. P., & Rona, P. A. (1985). Hydrothermal models for the generation of massive sulfide ore deposits. *Journal of Geophysical Research*, 90(B10), 8769–8783. doi:10.1029/JB090iB10p08769
- McCaig, A. M., Cliff, R. A., & Escartín, J. (2007). Oceanic detachment faults focus very large volumes of black smoker fluids. *Geology*, 35(10), 935–938.
- Ranero, C. R., & Reston, T. J. (1999). Detachment faulting at ocean core complexes. *Geology*, 27(11), 983–986.
- Reston, T. J., & McDermott, K. G. (2011). Successive detachment faults and mantle unroofing at magma-poor rifted margins. *Geology*, 39(11), 1071–1074.
- Spencer, J. E., & Welty, J. W. (1986). Possible controls of base-and precious-metal mineralization associated with Tertiary detachment faults in the lower Colorado River trough, Arizona and California. *Geology*, 14, 195–198.
- Tucholke, B. E., Behn, M. D., & Buck, W. R. (2008). Role of melt supply in oceanic detachment faulting and formation of megamullions. *Geology*, 36(6), 455–458.
- Tucholke, B. E., Lin, J., & Kleinrock, M. C. (1998). Megamullions and mullion structure defining oceanic metamorphic core complexes on the Mid-Atlantic Ridge. *Journal of Geophysical Research*, 103(B5), 9857–9866. doi:10.1029/98JB00167

Models and Timing of Core Formation

Sean P. Funk

Over the past one hundred years, important discoveries about the core have come from seismology, magnetism, and geodesy. However, the exact mechanisms involved in the formation of the core are still a mystery. Important constraints on core formation can be placed from Hf-W isotopes [1] and looking at siderophile (metal-loving) elements in the mantle [2]. Hostetler and Drake^[3] proposed that a "magma ocean", a consequence of accretion, formed on Earth. The "homogeneous magma ocean hypothesis" envisions that shock-induced melting caused metal and silicate to segregate, with the metal sinking toward the center of the proto-Earth.

Despite the initial success of the model in predicting the large-scale general distribution of metals and silicates, it fails to explain certain aspects of trace element geochemistry and isotope systematics. For example, the "excess siderophile element anomaly" (ESEA) in particular has been very troublesome to explain[4]. The ESEA demonstrates that the mantle, although depleted in siderophile elements relative to chondrites, is greater than predicted with known high-pressure equilibrium metal-silicate partitioning coefficients[5].

In this review, I will discuss three alternative hypotheses on core formation, and evaluate and critique each model with respect to the chemistry and physics involved. The first is core-core disequilibrium mixing, whereby the cores of differentiated bodies merge together quickly[6]. Here, the physics of emulsification become important to evaluate. Descending metal droplets are subject to Rayleigh-Taylor and Kelvin-Helmholtz instabilities, which act to tear the droplets apart. Only the largest descending cores may survive hybridization. Another is "inefficient core formation", where metallic material gets trapped within the mantle, later to be re-oxidized and redistributed[7]. The mode of metal transport, percolation versus dyking or diapirism, become important. Based on the dihedral angle (θ) of Earth materials at high-pressure, it seems unlikely that this is a viable model. The last is known as the heterogeneous accretion (or late veneer) model, where the composition of the accreting material changes with time[8]. During the late stages of core formation, a "late veneer" of material added siderophile elements into the mantle[8]. At present, this is the most widely accepted hypothesis that best explains core formation.

References

- [1] Kleine, T., Münker, C., Mezger, K. And Palme, H. (2002) Rapid accretion and early core formation on asteroids and the terrestrial planets from Hf-W chronometry. *Nature* **418**: 952-955
- [2] Meisel, T., Walker, R. J., and Morgan, J. W. (1996) The osmium isotopic composition of the Earth's primitive upper mantle. *Nature* **383**: 517–520.
- [3] Hostetler, C.J. and Drake, M.J. (1980) On the early global melting of the terrestrial planets. *Proceeding of the 11th Lunar and Planetary Science Conference*, pp. 1915-1929.
- [4] Walter, M.J., Newsom, H.E., Ertel, W., and Holzheid, A. (2000) Siderophile elements in the Earth and Moon: Metal/silicate partitioning and implications for core formation. In: Canup, R.M. and Righter, K. (eds) *Origin of the Earth and Moon*, pp. 265-290. Tucson, AZ: University of Arizona Press.
- [5] Li, J, and Agee, C.B. (1996) Geochemistry of mantle-core differentiation at high-pressure. *Nature* **381**: 686-689.
- [6] Halliday, A.N. (2006) The origin of the Earth: What's New? *Elements* **2**: 205-210.
- [7] Jones, J.H., and Drake, M.J. (1986) Geochemical constraints on core formation in the Earth. *Nature* **322**: 221-228.
- [8] Wänke, H (1981) Constitution of terrestrial planets. *Philosophical Transactions of the Royal Society* **303**: 287-302.

Monday November 26, 9:30 am in BGS 1053

Dominant gliding versus pure spreading in passive margins: the effect of differential sedimentation on initiating salt tectonics

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Dominant gliding and pure spreading are both gravity driven salt tectonics models, but they form in two different ways. Dominant gliding is primarily gravity driven via marginal tilt with a minor implication for the impact of sedimentation. Conversely, pure spreading is driven only by differential sediment loading in a horizontal environment. When discussing the application of the dominant glide model as opposed to the pure spreading model in passive margins, determining the effect of differential sediment loading on initiating salt tectonics is a major point of contention.

In “Salt tectonics at passive margins: Geology versus models” by Brun and Fort, 2011, the application of numerical and experimental models concludes that in order for differential sedimentary loading to initiate salt tectonics, abnormally thick and/or dense sediments, in conjunction with anomalously low sediment friction would be required. According to Brun and Fort, 2011, dominant gliding is more efficient and significantly easier to initiate than pure spreading. These findings, specifically 1) the equations used, 2) the laboratory experiments, and 3) the geological analogies, were hotly contested by Rowan in a discussion article published in 2012. Rowan advocated that pure spreading has a much larger impact than implied by Brun and Fort, 2011. However, the majority of the concerns by Rowan, 2012, are based on 1) numerical models that don’t account for the same variables, 2) laboratory experiments done by other authors working with Rowan, and 3) an overall misunderstanding and misrepresentation of what was published by Brun and Fort, 2011. Despite the opposing views, both authors agree that the effectiveness of models in a geologic setting cannot be guaranteed and that both dominant gliding and pure spreading are normally occurring processes.

References

- Brun, J.-P. and Fort, X. 2011. Salt tectonics at passive margins: Geology versus models (Review Article). *Marine and Petroleum Geology*. vol 28. p 1123-1145.
- Brun, J.-P. and Fort, X. 2011. Salt tectonics at passive margins: Geology versus models – Reply (Discussion). *Marine and Petroleum Geology*. vol 37. p 195-208.
- Brun, J.-P., Mauduit, T., 2008. Origin of rollovers in salt tectonics: the inadequacy of the listric fault model. *Tectonophysics*. vol 457. p 1-11.
- Brun, J.-P., Mauduit, T., 2009. Salt rollers: structure and kinematics from analogue modelling. *Marine and Petroleum Geology*. vol 26. p 249 - 258.
- Davison, I., Anderson, L., Nutall, P., 2011. Salt deposition, loading and gravity drainage in the South Brazilian salt basin. In: *Salt Tectonics, Sediments and Prospectivity*. Geological Society, London, Special Publications.
- Gaullier, V., Vendeville, B.C., 2005. Salt tectonics driven by sediment progradation. Part II: Radial spreading of sedimentary lobes prograding above salt. *American Association of Petroleum Geologists Bulletin*. vol 89. p 1081 - 1089.
- Gemmer, L., Ings, S.J., Medvedev, S., Beaumont, C., 2004. Salt tectonics driven by differential sediment loading: stability analysis and finite element experiments. *Basin Research*. vol 16. p199 - 219.
- Gemmer, L., Beaumont, C., Ings, S.J., 2005. Dynamic modelling of passive margin salt tectonics. effects of water loading, sediment properties, and sedimentation patterns. *Basin Research*. vol 17. p 383 - 402.
- Hudec, M.R., Jackson, M.P.A., 2007. Terra infirma: understanding salt tectonics. *Earth Science Reviews*. vol 82. p 1 - 28.
- Rowan, M.G., Peel, F.J., Vendeville, B.C., Gaullier, V., 2012. Salt tectonics at passive margins: geology versus models - discussion. *Marine and Petroleum Geology*. vol 37. P 184 - 194.
- Schultz-Ela, D.D., 2001. Excursus on gravity gliding and gravity spreading. *Journal of Structural Geology*. vol 23. p 725 - 731.
- Vendeville, B.C., 2005. Salt tectonics driven by sediment progradation: Part I: Mechanics and kinematics. *American Association of Petroleum Geologists Bulletin*. vol 89. p 1071 - 1079.
- Waltham, D., 1996. Why does salt start to move? *Tectonophysics*. vol 282.

Weathering-induced metal-enrichment processes, the aluminium and nickel cases

Martin Arce

This paper reviews the processes responsible for the enrichment of metals through intense weathering of rocks. It provides a basis for understanding the general weathering phenomena and its relationship with tectonic and climate. It also describes the main characteristics of the two major weathering-related deposits in terms of mine production, aluminium and nickel laterites, as well as their basic concentration mechanisms. Weathering results of the interaction between the hydrosphere, biosphere and lithosphere under determined tectonic, climatic and topographic conditions, where some elements of primary minerals are lixiviated and secondary minerals are produced as residua. Laterite is the most generally accepted denomination for the product of intense weathering in humid, cold to warm climates and sub-artic to tropical regions during sufficient time under conditions of tectonic stability. The final product of weathering is a mineral assemblage of the least soluble minerals and the most resistant primary minerals. The evolution of laterites on parent rocks with pre-concentration of certain metals cause rock destruction and reorganisation of these elements in new supergene associations which either stay within the profile or migrate in solution.

Bauxite is the only source used for production of alumina on industrial scale, world production during 2011 reached 211000 thousand metric dry tons. Bauxite is a lateritic rock characterized by the extreme enrichment of aluminum hydroxide minerals, such as gibbsite, boehmite and diaspore, together with iron oxides, kaolinite and less anatase. Laterites and bauxites are generated in tropical environments by intense weathering with the consequent enrichment in iron (laterites) and in alumina (bauxites) to ore grades. If compared to laterites, bauxites are generated by stronger leaching. Dissolved silica concentration is lowered by the intense leaching, enhancing formation of gibbsite instead of kaolinite. A classification based on the tectonic frame includes bauxites in uplift areas, in subsiding platforms, and in carbonate platforms.

Nickeliferous laterites currently represent the major nickel reserves, approximately 48,000,000 metric tons of nickel content. Even though 60% of world nickel reserves are composed by lateritic nickel, only 40% of world nickel production is generated from this source. It is due to the difficulties and the higher energy consume in metallurgical processing of nickel oxides compared to nickel sulfides. Progressive and intensive weathering of ultramafic rocks generally under tropical conditions (it also occurs in wet cold regions at much lower rates) generates economic concentrations of nickel, platinum group elements and chromium. Cobalt and copper are usual by-products but also can constitute deposits by themselves. Based on its mineralogy, lateritic nickel deposits are classified in hydrous silicate deposits, clay silicate deposits and oxide deposits. The importance of weathering-forming deposits processes and the convenience and/or need of mining them are highlighted and demonstrated in this contribution through the examples of the two major commodities mined from them.

References

- Anand, R.R. and Butt, C.R.M. 2010. A guide for mineral exploration through the regolith in the Yilgarn craton, Western Australia. *Australian Journal of Earth Sciences*, **57**: 1015-1114.
- Elias, M. 2004. Nickel laterite deposits – geological overview, resources and exploitation. In: *Giant ore deposits: characteristics, genesis and exploration*. Edited by DR Cooke and J. Pongratz. CODES Special Publication 4, Centre for Ore Deposit Research, University of Tasmania, 205-220.
- Freyssinet, P., Butt, C., Morris, R. and Piantone, P., 2005. Ore-forming processes related to laterite weathering. *Economic Geology* 100th Anniversary volume, p. 681–721.
- Glazkovsky, A., Gorbunov, G., and Sysoev, F.A. 1977. Deposits of nickel. In Smirnov, V.I., ed., *Ore deposits of the USSR, Volume II*: London, Pitman Publishing, p. 3–79.
- Gleeson, S.A., Butt, C.R.M., and Elias, M. 2003. Nickel laterites: a review. *SEG Newsletter, Society of Economic Geologists*, **54**: 1;12-18.
- Jadhav, G.N., Sharma, N., and Sen, P. 2012. Characterization of bauxite deposits from Kachchh area, Gujarat. *Journal of the Geological Society of India*, **80**: 351-362.
- Mudd, G. 2009. Nickel sulfide versus laterite: the hard sustainability challenge remains. In: Proc. “48th Annual Conference of Metallurgists”, Canadian Metallurgical Society, Sudbury, Ontario, Canada, August 2009.
- Ndjigui, P. and Bilong, P. 2010. Platinum-group elements in the serpentinite lateritic mantles of the Kongo-Nkamouna ultramafic massif (Lomie region, South-East Cameroon). *Journal of Geochemical Exploration*, **107**: 63-76.
- Schellmann, W., 1981. Considerations on the definition and classification of laterites. *Proceedings of the International Seminar on Lateritisation Processes*, Trivandrum, India. A.A. Balkema, Rotterdam, p. 1-10.
- Schellmann, W. 1994. Geochemical differentiation in laterite and bauxite formation. *Catena*, **21**: 131-143.
- Taylor, G., Eggleton, R.A., Holzhauer, C.C., Maconachie, L.A., Gordon, M., Brown, M.C., and McQueen, K.G. 1992. Cool climate lateritic and bauxitic weathering. *Journal of Geology*, **100**: 669-677.
- Valeton, I. 1994. Element concentration and formation of ore-deposits by weathering. *Catena*, **21**: 99-129.
- Yamaguchi, Kosei E. 2010. Iron isotope compositions of Fe-oxide as a measure of water-rock interaction: An example from Precambrian tropical laterite in Botswana (Report). **2**. p. 3.
- Bray, E.L., 2012. Bauxite and alumina. *Commodity Statistics and Information*. U.S.G.S. Available from: <http://minerals.usgs.gov/minerals/pubs/commodity/bauxite/mcs-2011-bauxi.pdf> [Accessed 5 November 2012].

Geophysical Techniques for shallow Subsurface Ground Penetrating Radar (GPR), Multi-Channel Analysis of Shear Waves (MASW) Seismic Refraction and Reflection

Wajahat Ali

Shallow subsurface is an important geological zone which is directly related to human life in terms of water supply, agriculture and ecosystems. Several geophysical techniques are available for the investigation and characterization of shallow subsurface. Most commonly used are the electrical methods, seismic refraction and reflection methods, multichannel analysis of surface waves, gravity, magnetic, electromagnetic induction and ground penetrating radar. Each of the techniques is based on a specific physical law. An attempt has been made in this paper to describe the seismic and ground penetrating radar techniques with their advantages and limitations.

Seismic reflection and refraction methods involves the study of body waves (P and S waves) travelling through the earth interior, reflecting and refracting on the interfaces and discontinuities having different acoustic impedances. Multichannel analysis of surface waves involves the study of surface waves travelling along the air and earth interface suffering dispersion, whereas ground penetrating radar involves the analysis of reflected electromagnetic signals from the objects having different dielectric constants. Reflection method describes the subsurface stratigraphy and discontinuities (such as faults and erosional surfaces) efficiently. Refraction method is mostly to estimate the depth to the bedrock. Analysis of dispersive surface waves identifies the zones having voids, weathered and fractured bedrock. Ground penetrating radar uses the principle of reflection of electromagnetic energy which identifies subsurface utilities, fractures, void and archeological sites.

Each technique has its own limitations and drawbacks. The depth range of ground penetrating radar depends upon the electrical conductivity of the subsurface. In ground having high salt content, the depth of penetration may not reach few centimeters where as in ice it may reach several hundred meters. On the other hand a dipping layer in shallow subsurface significantly affects the shear wave inversion results apart from the maximum resolvable wavelength of the fundamental mode with respect to the spread length. Reflection and refraction methods will depend upon the spread length, offsets and kind and frequency contents of the source used.

References

- Ali, H., Atef, Uchenna, H., Aboajal, Gaunt, D., Kelly, H., Liu1, and Stephen, S., Gao. 2009. Geophysical Investigation of the Shallow Subsurface in St. Charles and St. Louis Counties, Missouri. American Association of Petroleum Geologist Article No. 9011
- Annan, A.P., and Davis, J.L. 1997. Ground penetrating radar coming of age at last. *In* proceeding of the Fourth Decennial International Conference on Mineral Exploration, Electrical and Magnetic Methods, Paper 66, pp. 1512-1522
- Brain, J., Moorman. 2001. Ground penetrating radar applications in paleolimnology. Available from University of Calgary, Calgary pp. 2
- Choon, B., Park, Richard, D., Miller, Xia, J., and Ivanov, J. 2007. Multichannel analysis of surface waves (MASW)—active and passive methods. *Leading edge journal*
- Choon, B., Park, Richard, D., Miller and Xia, J. 1997. Multi-channel analysis of surface waves. Open file Report No. 97-10, Kansas Geological Survey. pp. 10-12
- Choon, B., Park, Richard D., Miller, and Xia, J., 1999. Multi-channel analysis of surface waves. *Journals of society of exploration geophysics*. VOL. 64, NO. 3. pp. 800-808
- Donald, B., Hoover, Douglas, P., Klein, and David, C., Campbell. 1997. Geophysical Methods in Exploration and Mineral environmental investigation. Personal communication
- Dojack, L. 2012. Ground penetrating radar theory, data collection, processing and interpretation. Available from University of British Columbia, British Columbia pp. 2-11
- Finck, F. 2003. Introduction of a ground penetrating radar system for investigations on concrete structures. *Otto-Graf Journal* Vol. 14, pp.36
- Gosar A., Stopar, R., and Roser, J. 2008. Comparative test of active and passive multichannel analysis of surface waves (MASW) methods and microtremor HVSr method. *Scientific paper journal* Vol. 55, No. 1, pp. 41-66
- Milsom, J. Eriksen, A. 2011. *Field geophysics*, 4th edition. Jhon Wiley and Sons, Ltd., Virginia, USA
- Remke, L., Van Dam. 2001. Causes of ground-penetrating radar reflections in sediment. Vrije University, Netherlands
- Edwin S. Robinson. 1988. *Basic exploration geophysics*. Jhon Wiley and Sons, Ltd., West Sussex, UK
- Richard D., Miller, Xia J., Choon B., Park, and Julian M., Ivanov. 1999. Multichannel analysis of surface waves to map bedrock. *Leading edge journal*. pp. 1392-1393
- Richard, J., Yelf. 2007. Application of Ground Penetrating Radar to Civil and Geotechnical Engineering. *Electromagnetic Phenomena*, Vol.7, No.1. pp. 103-105

Database for flooding susceptibility, hazard, and vulnerability assessment

Hadis Samadi Alinia

Disasters are natural and human-caused events that have an adverse impact on a community, region, or nation. Events associated with a disaster can overwhelm response resources and have damaging economic, social, or environmental impacts. Floods are some of the most common and costly natural disasters around the world. These events occur in most countries, and cause the most deaths and it is expected that the extent of flooding increase under the influence of climate change and economic development.

There is a need to clarify the nature and impacts of the flood hazards in a hazard analysis process. The assumption is that negative effects of disasters can be reduced through preparation. Much effort has gone into preparing people of disaster prone areas to withstand the effects of disasters. A number of approaches have been applied for disaster preparedness. They focused on prior information regarding impending disasters, frequency and severity of disasters, causes, effects of disasters and their reduction or mitigation, perception of risk, removing people from disaster prone areas, disaster preparedness, coping and adjustment, post-disaster rebuilding and return to normalcy.

Among these, Geo-information and remote sensing are proper tools to enhance functional strategies for increasing awareness on natural hazard prevention and for supporting research and operational activities devoted to disaster reduction. GIS along with remote sensing has become the key tool to delineate of flood prone areas and development of flood hazard maps indicating the risk areas likely to be inundated by significant flooding along with the damageable objects maps for the flood susceptible areas. Producing rescue and flood vulnerability map and updating it using satellite images would help for evacuation and dispatching resources and aid scared people to the safe regions in a short time.

As producing susceptibility, hazard, and vulnerability maps is composed of various criteria involved in the flood disaster, reliable, up-to-date, and accurate geospatial and non-spatial data is significant. With respect to the environmental factors used in Flooding hazard assessment, there is a tendency to utilize those data that are easily obtainable from Digital Elevation Models and satellite imagery, whereas less emphasis is on those that require detailed field investigations.

This paper is the review on the types of spatial and non-spatial data needed in this hazard case and the approaches for obtaining them and also mapping the infected zones. This paper is a review in collecting spatial and non-spatial information on environmental factors with a focus on Digital Elevation Models, geology and soils, geomorphology, land use and elements at risk.

References

- Ajmar A, Perez F, Terzo O. 2008. *WFP spatial data infrastructure (SDI) implementation in support of emergency management*. XXI congress of the International Society for Photogrammetry and Remote Sensing.
- Alvarado-Aguilar, D., Jimenez, J. A., Nicholls, R.J.. 2012. *Flood hazard and damage assessment in the Ebro Delta (NW Mediterranean) to relative sea level rise*. 62:1301–1321.
- Baldassarre, G.D, Schumann, G, Bates, p. 2009. *Near real time satellite imagery to support and verify timely flood modeling*, 23(5): 799-803.
- Giardino, M., Perotti, L., Lanfranco, M., Perrone, G. 2012. *GIS and Geomatics for disaster management and emergency relief: a proactive response to natural hazards*, Applied Geomatics, 4: 33-46.
- IPCC (Intergovernmental Panel on Climate Change). 2007. *The physical science basis. contribution of working group I to the fourth assessment report of the intergovernmental panel on climate change*. Cambridge University Press, Cambridge.
- Jabed Abdul Naser, Md., Dushmanta D. 2011. *Analysis of flood vulnerability and assessment of the impacts in coastal zones of Bangladesh due to potential sea-level rise*, Natural Hazards, 61:729–743.
- Lawal, D.U, Matori. A.N, Hashim, A.M., Chandio, I.A., Sabri, S., Balogun, A.L, Abba, H.A., 2011. *Geographic Information System and Remote Sensing Applications in Flood Hazards Management: Review*. Applied Sciences, Engineering and Technology. 3(9): 933-947.
- Noji, E. K., and Lee, C. Y. 2005. *Disaster preparedness*. In H. Frumkin (Ed.), *Environmental health: From global to local*, San Francisco, CA: Jossey-Bass, pp. 745–780.
- Pine, J. C. 2008. *Natural hazards analysis: reducing the impact of disasters Boca*. Auerbach Publications.
- Kaizhong Li • Shaohong Wu • Erfu Dai • Zhongchun Xu. 2012. *Flood loss analysis and quantitative risk assessment in China*. Natural Hazards. 63:737–760.
- Wood, M., Kovacs, D., Bostrom, A., Bridges, T., Linkov, Igor, 2012. *Flood Risk Management: US Army Corps of Engineers and Layperson Perceptions*. Risk Analysis. 32: 1349-1368
- Pradhan, B. 2009. *Flood susceptible mapping and risk area delineation using logistic regression, GIS and remote sensing*. Spatial Hydrology. 9(2): 1-18.
- Thompson A., Clayton J. 2002. *The role of geomorphology in flood risk assessment*. P I Civil Eng-Civ En. 150:25–29.
- Miceli, R., Sotgiu, I., & Settanni, M. 2008. *Disaster preparedness and perception of flood risk: a study in an alpine valley in Italy*. Environmental Psychology. 28, 164–173.