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# Landslide processes in the south-central Mackenzie River valley region, Northwest Territories

David Huntley and Alejandra Duk-Rodkin

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**Abstract:** As part of the Northern Energy Development Mackenzie Valley Project, an inventory of terrain, landforms and geomorphic processes is currently being compiled for the Camsell Bend (NTS 95 J), Root River (NTS 95 K), Dahadinni River (NTS 95 N), and Wrigley (NTS 95-O) 1:250 000 scale map sheets. This database will provide insight into the types of landslides and their geological setting, dimensions, age, distribution, activity, and density in the study area. The landslide data recorded by this study will provide calibration for future predictive mapping and hazard analyses in the Mackenzie River valley transportation corridor.

**Résumé :** Dans le cadre du projet de la vallée du Mackenzie visant le développement de l'énergie dans le Nord, nous procédons actuellement à la compilation d'un inventaire du terrain, des formes de relief et des processus géomorphologiques dans les régions visées par les cartes suivantes à l'échelle du 1/250 000 : Camsell Bend (SNRC 95 J), Root River (SNRC 95 K), Dahadinni River (SNRC 95 N) et Wrigley (SNRC 95-O). Cette base de données fournira un aperçu des types de glissements de terrain dans la région à l'étude, de leur contexte géologique, de leurs dimensions, de leur âge, de leur répartition, de leur activité et de leur densité. Les données sur les glissements de terrain consignées dans le cadre de cette étude serviront de données d'étalonnage pour les futures activités de cartographie prévisionnelle et d'analyse des risques dans la voie de communication de la vallée du Mackenzie.

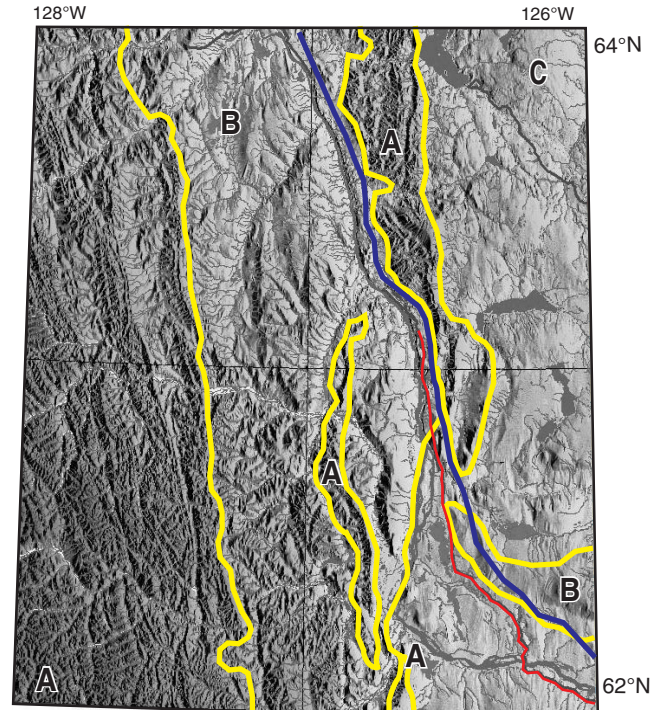
## INTRODUCTION

Geological and geomorphological studies in the Mackenzie River valley region by the Geological Survey of Canada (GSC) provide valuable information on bedrock, surficial deposits, permafrost, and landslides for the Northern Energy Development Mackenzie Valley Project. These data are essential for effective land management decisions regarding pipeline, highway, and settlement construction, and the extraction of aggregate and groundwater resources in the region (Fig. 1). Current research (e.g. DiLabio, 2005; Couture et al., 2005; Huntley and Duk-Rodkin, 2005; Pyle et al., 2005; Sidwell et al., 2005; Smith et al., 2005; Wang et al., 2005) builds on legacy terrain data (e.g. Aylsworth et al., 2000a, b; Dyke, 2000; Heginbottom, 2000; Aylsworth and Traynor, 2001; Duk-Rodkin and Hood, 2004), and will be applicable to a wide range of end users. The focus in this paper is on landslides as geoindicators of slope instability in the Mackenzie River valley transportation corridor and adjacent areas.

### Regional physiographic setting

During the 2005 field season, surficial deposits and landslides were described and mapped in the Camsell Bend (NTS 95 J), Root River (NTS 95 K), Dahadinni River (NTS 95 N), and Wrigley (NTS 95-O) 1:250 000 NTS map sheets. The study area comprises three major physiographic elements (Fig. 2):

A) Mackenzie Mountains and McConnell Range (Franklin Mountains): montane terrain with exposed folded and thrust-faulted carbonate and clastic sedimentary rocks, characterized by high relief and steep slopes (Fig. 3).



**Figure 2.** Generalized physiography of the south-central Mackenzie River valley region, Northwest Territories: **A)** montane terrain, **B)** foothills and intermontane plateaus, **C)** plains (*modified from Duk-Rodkin and Lemmen, 2000*). Symbols: red line, Mackenzie Highway; blue line, proposed pipeline route (AMEC Americas Limited, 2005).

B) Mackenzie Lowland: drift-covered foothills and intermontane plateaus, underlain by folded and faulted weak shale, limestone, and sandstone, and deeply incised by tributaries draining to the Mackenzie River (Fig. 3).

C) Great Slave Plain: till and glaciolacustrine plains blanketed by outwash and extensive organic deposits, and overlying relatively undeformed shale, sandstone, and limestone (Fig. 3).

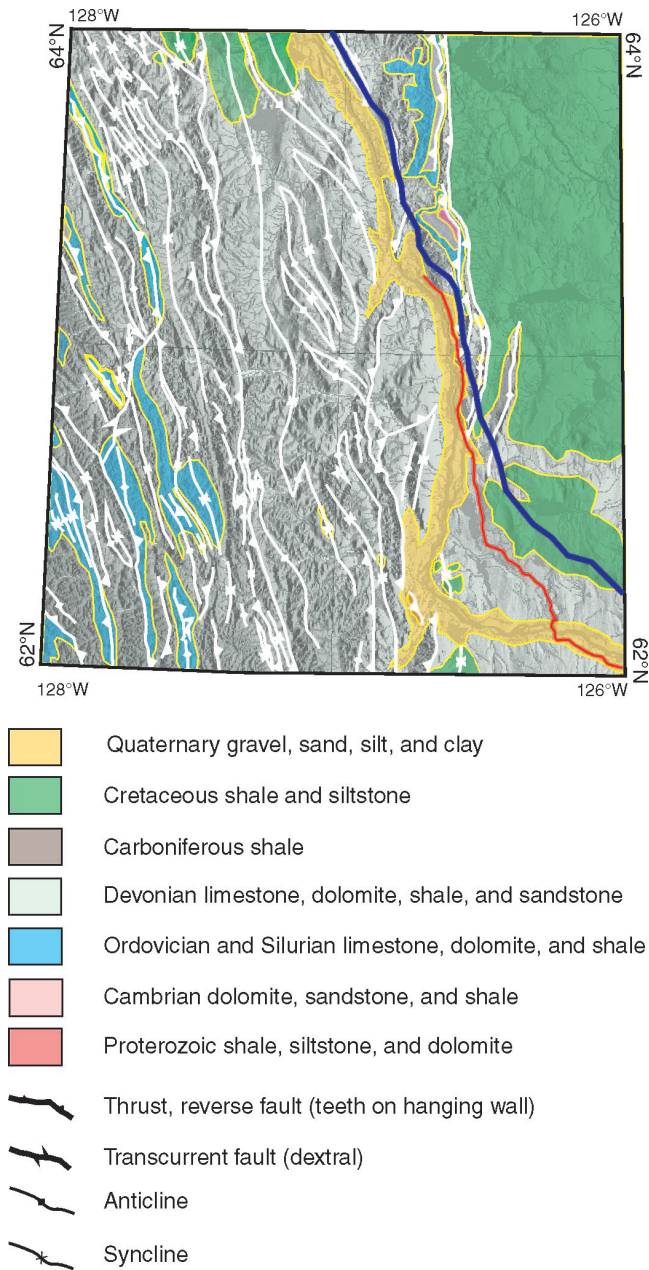
## TERRAIN AND LANDSLIDE MAPPING

To better understand landslide processes and their impacts, an important terrain-mapping objective was to identify and classify individual and groups of landslides. Surficial geology and applied terrain maps to be produced as part of this study will provide an objective quantitative inventory of landslide hazards, showing the range, distribution, and magnitude of events by the number and size of mass movements (cf. Aylsworth and Traynor, 2001; Duk-Rodkin and Hood, 2004). These maps can be applied at a later date to more elaborate landslide-hazard analyses, including the spatial probability of occurrence, magnitude, and intensity of landslides in the initiation zone, and the extent and nature of landslide run-out. Landslide risks can also be derived from the



**Figure 1.** Limits of the Northern Energy Development Mackenzie Valley Project and location of the 2005 study area.

landslide-hazard maps by including elements at risk, vulnerability, and consequence in the Mackenzie River valley transportation corridor.



**Figure 3.** Simplified geology of the south-central Mackenzie River valley region, Northwest Territories (*modified after* Douglas and Norris, 1976a–d; Okulitch, work in progress, 2006), showing the Mackenzie Highway (red line) and proposed pipeline route (blue line; AMEC Americas Limited, 2005). See Tables 1 to 3 for explanation of terrain-classification symbols.

## Terrain and landslide classification

Surficial units, landforms, and geomorphic processes were pre-typed onto 1:40 000 to 1:60 000 scale black-and-white stereopair airphotos (from 1947, 1961, and 1971). Selected polygons were benchmarked during the 2005 field season. At more than 800 ground-control stations, surficial deposits were defined on the basis of landform associations, texture, sorting, fissility, colour, sedimentary structures, consolidation, presence of permafrost, and contact relationships. After revising the pre-typed airphotos, polygon linework and onsite symbols were digitized in ArcMap and terrain labels transcribed into a corresponding database. When combined with georeferenced field data, the resulting geographic information system will be used to generate a range of applied terrain products, including landslide-hazard maps.

Landslide events and terrain modified by mass wasting were classified according to type of earth material involved (e.g. bedrock or surficial deposits), rate of movement, presence of permafrost, nature of movement and associated processes (e.g. fall, topple, flow, slide, spreading), status of activity (i.e. active or stable), triggering mechanism(s), aspect, and dimensions. Other data captured during landslide pre-typing included information on bedrock and geological structures (strike directions of faults and fold axes). Landform, surficial-unit, and geomorphic-process codes (Tables 1–3) used in this study are derived from

**Table 1.** Earth-material classification for the south-central Mackenzie River valley region, Northwest Territories.

Map symbol	Earth material
<b>A</b>	Alluvial
<b>C</b>	Colluvial
<b>E</b>	Eolian
<b>G</b>	Glaciofluvial
<b>L</b>	Glaciolacustrine
<b>O</b>	Organic
<b>R</b>	Bedrock
<b>T</b>	Moraine (till)
<b>U</b>	Undifferentiated

**Table 2.** Surface-expression classification for the south-central Mackenzie River valley region, Northwest Territories. See Table 4 for hazard rating.

Map symbol	Surface expression	Description
<b>Sediment depth</b>		
<b>b</b>	Blanket (>2 m)	
<b>v</b>	Veneer (<2 m)	
<b>Surface forms</b>		
<b>p</b>	Plain (<5%)	Planar surface
<b>t</b>	Terrace	Single/stepped platform
<b>(G) d</b>	Delta	Fan to irregular platform
<b>(T) d</b>	Drumlin	Elongated hillock >1 m
<b>f</b>	Fan (<26%)	Cone shape
<b>h</b>	Hummock (26–70%)	Local relief >1 m
<b>r</b>	Ridge (26–70%)	Elongated hillocks >1 m
<b>x</b>	Complex	Complex of surface forms

mapping conventions used by the Geological Survey of Canada (e.g. Evans and Savigny, 1994; Aylsworth and Traynor, 2001), the Mackenzie Gas Project Landform Sensitivity Atlas (AMEC Americas Limited, 2005), and terrain analysts in British Columbia and the Yukon (e.g. Varnes, 1978; Cruden and Varnes, 1996; Resource Inventory Committee, 1996; Howes and Kenk, 1997; Huscroft et al., 2004). This classification scheme is specific to the study area and accommodates complex events involving a range of materials and flow regimes over several orders of magnitude and time.

**Table 3.** Geomorphic-process classification for the south-central Mackenzie River valley region, Northwest Territories. See Figure 4 for classification of landslides and periglacial processes.

Process type	Map symbol	Geomorphic process
Fluvial	B	Braiding channel
	M	Meandering channel
Erosional	D	Deflation
	K	Karst processes
	V	Gully erosion
	W	Washing
Mass movement	F	Slow mass movements
	R	Rapid mass movements
Glacial	E	Channeling by meltwater
	H	Kettled
Periglacial	X	Permafrost processes

**Table 4.** Terrain-stability classification for the south-central Mackenzie River valley region, Northwest Territories.

Slope class	Terrain-stability class definition <i>Following disturbance of slope by natural events and human modification</i>
I	No significant slope-stability problem
II	Low likelihood of slope failure
III	Moderate likelihood of slope failure
IV	High likelihood of slope failure
V	Very high likelihood of slope failure

**Table 5.** Terrain and geomorphic-process associations for the south-central Mackenzie River valley region, Northwest Territories.

Terrain unit	Geomorphic process(es)	Ice content	Drainage	Terrain stability rating
pO, fO	X	20–80 %	Poor	I to V
Ev, Eb, Er	D	Low	Good	I to II
Ax, Ap, Af, Ar	B, M	Low	Good	III to V
Cx, Cz, Cy, Cf, Cb, Cv	F, R, X	Variable	Good	III to V
Lx, Lp, Lt, Lb, Lv	F, R, V, X	10–25 %	Poor	I to V
Gx, Gp, Gt, Gd, Gf	D, F, R, V, X	Low	Good	II to IV
Tx, Td, Tr, Tb, Tv	E, F, R, X	10–25%	Moderate	I to III

## Subjective geomorphic analysis

Qualitative terrain-stability classes (Table 4) were assigned to each map polygon based on airphoto interpretation, field observations, previous studies, and experience of the terrain analyst. Terrain stability was defined in terms of predictable changes in terrain and the probability of occurrence of landslides following disturbance (either natural or anthropogenic). Classes were assigned subjectively by the judgmental weighting of the following key terrain attributes observed in a polygon:

- surficial materials
- geomorphic processes, active and inactive
- slope gradient and local relief
- factors such as soil drainage, soil depth, vegetation cover, and fire history
- presence of existing landslides

Terrain-stability classes will be coloured on landslide-hazard maps to distinguish zones of extreme and high landslide hazard (red), zones of moderate hazard (yellow), and zones of little or no perceived hazard (green).

## LANDSLIDE ANALYSIS AND GEOMORPHIC MODEL

The interaction between high relief, steep slopes, heavy precipitation, complex tectonic and glacial history, and the nature of surficial deposits and bedrock has produced a variety of landslides along the Mackenzie River valley transportation corridor (Aylsworth and Traynor, 2001; Duk-Rodkin and Hood, 2004; Huntley and Duk-Rodkin, 2005; Couture et al., 2005). Table 5 summarizes the dominant terrain units and geomorphic processes observed in the study area during 2005. Estimates of ice content and nature of drainage are based on field observations and data from adjacent areas (Aylsworth et al., 2000a, b; Dyke, 2000; Heginbottom, 2000). Slope classes were derived from contour intervals and ground measurement at field stations.

The geomorphic model (Fig. 4) developed as part of this study is intended to provide baseline geomorphic information for a range of potential end-users, including research scientists, construction engineers, land-use managers, terrestrial ecologists, and information specialists. The distribution, activity, and density of landslides recorded as part of current research will help calibrate future predictive mapping and hazard analyses in the Mackenzie River valley transportation corridor. Throughout the study area, bedrock and debris slopes exceeding 28% and 30 m in height are potentially unstable. In the Mackenzie Mountains and foothills of the Mackenzie Lowland, rock falls and topples are triggered by gravitational spreading, seismic activity, and weathering processes. Some mountain summits display cryoplanation

## TERRAIN CLASSIFICATION and LANDSLIDE CLASSES

### Rapid mass movement (R)

- Rb Rockfall
- Rd Debris flow
- Rf Debris fall
- Rm Bedrock slide (slump)
- Rt Bedrock topple
- Ru Debris slide (slump)
- Rx Landslide complex

### Slow mass movement (F)

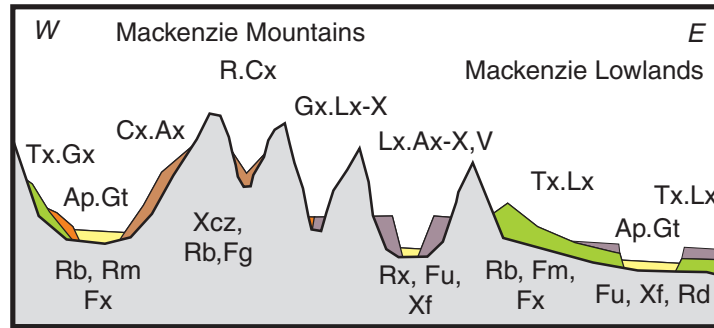
- Fg Gravitational spreading
- Fm Bedrock slide (slump)
- Fu Debris slide (slump)
- Fx Landslide complex

### Permafrost (X)

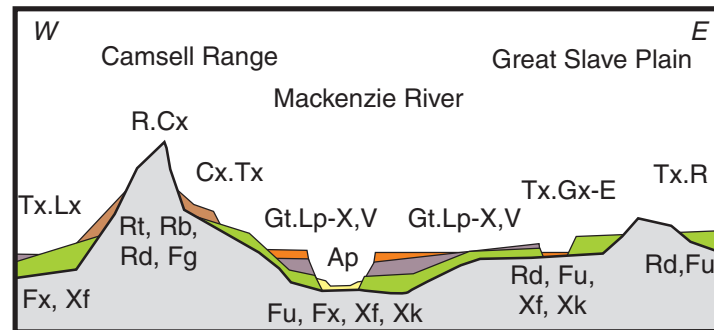
- Xc Cryoplanation
- Xe Thermal erosion by water
- Xf Retrogressive thaw flow slides
- Xk Thermokarst processes
- Xz Gelifluction

### Other Processes

- E Meltwater channel
- V Gully erosion



Root River (NTS 95 K) and Dahadinni River (NTS 95 N)



Camsell Bend (NTS 95 J) and Wrigley (NTS 95-O)

Cross-sections are conceptual and not to scale

**Figure 4.** Geomorphic model for the south-central Mackenzie River valley region, Northwest Territories. See Tables 1 to 3 for explanation of terrain classification.

surfaces on which weathered bedrock and surficial deposits have been locally remobilized by gelifluction (Fig. 4). In the Mackenzie Lowlands and over the western Great Slave Plain, unfrozen and saturated shale beds can move as saturated debris flows or rock slides. The Devonian Fort Simpson Formation is the rock unit most susceptible to failure. Most landslides, however, occur in glaciolacustrine sediments and fine-grained, ice-rich till. Because of the low shear strength of water-saturated sediments, thawing of ice-rich, fine-grained sediments on gentle slopes (5–27%) usually gives rise to active-layer detachments and retrogressive-thaw flows. Rapid debris flows in fine-grained deposits occur in areas with high relief or where saturated sand deposits occur (Fig. 4).

Along the Mackenzie River and its tributaries, frozen gravel and sand can move as blocks in rotational and translational debris slides, with failure occurring along planes developed in underlying weak, unfrozen, glaciolacustrine clay and silt. Where the height of channel banks and thickness of coarse-grained sediment are greater than the depth of permafrost, groundwater flow below the permafrost and above underlying impermeable deposits gives rise to artesian flow conditions (Aylsworth et al., 2000b; Dyke, 2000). Retrogressive-thaw flows and debris flows may develop if river erosion removes frozen, ice-cemented colluvium at the base

of slopes. Thermokarst subsidence and thermal erosion of frozen organic deposits, outwash, and lake sediments by streams will also trigger translational slides along valley floors (Fig. 4).

Dominant factors in triggering landslides and determining the speed of mass movement are the water content of bedrock and surficial material, and the occurrence of melting permafrost. Low water content reduces mass movement by increasing the cohesion of debris and shear strength. Saturating water levels decrease shear strength by increasing pore pressure, which triggers mass movement (Aylsworth et al., 2000b; Dyke, 2000; Couture et al., 2005; Wang et al., 2005). Other controls include distribution and intensity of annual taiga fires, earthquake activity, slope gradient, local relief, aspect of slope, debris thickness over bedrock, orientation of planes of weakness in bedrock, and bioclimatic factors (e.g. ice, precipitation, and vegetation cover).

For the current climate regime, there is a clear relationship between bank erosion and slope channel failure in the Mackenzie River valley (Aylsworth et al., 2000a, b; Dyke, 2000). Channel migration leads to overly steep and undermined riverbanks and terraces, and has triggered numerous complex landslides along the Mackenzie River and its tributaries (Huntley and Duk-Rodkin, 2005; Couture et al., 2005).

Removal of streamside vegetation and soil also leads to active-layer detachments, and retrogressive-thaw flows and slides in ice-rich terrain. Increased channel instability can be expected if there are changes in the amount and timing of precipitation, snow melt, and sediment supply to streams. Rates of bank erosion may increase as rivers adjust their gradients to accommodate changing discharge and sediment loads.

To reduce the possibility of landslides in the transportation corridor, construction should aim to minimally disturb steep slopes at stream and river crossings, and where gently sloping terrain dominated by fine-grained glacial sediment with discontinuous permafrost is indicated on surficial-geology and applied-terrain maps. Wherever possible, construction should employ passive techniques that avoid potentially unstable terrain identified by terrain mapping; installation of piles into permafrost; aggregate pads to insulate subjacent permafrost; and preservation of the protective cover. Nonpassive construction involving the removal of surficial cover to bedrock may be possible in areas where the drift cover is less than 5 m thick.

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